Biogeochemical Systems -- OCN 401

Readings: Schlesinger Chapter 7; Erwin (2009)

I. Redox Biogeochemistry in Aquatic Systems

II. Wetlands & Layered Microbial Habitats

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Aquatic redox

Oxidation-Reduction Chemistry Review

- Oxidation states, balancing equations
- Oxic-anoxic & redox potential
- Simple electrochemical cell

Redox reactions

- Chemical speciation
- Eh pH diagrams
- Redox reactions in nature

Biogeochemical reactions and their thermodynamic control

- Redox sequence of OM oxidation in aquatic environments
- Leading into wetlands, vertical profiles, time-series, and layered microbial habitats...

Wetlands & layers

Emphasize organic matter cycling

- e- donor, Storage
- Microbial metabolic pathways

Wetland habitats (& other layered microbial habitats)

- Types of wetlands
- Hydrology, soils,
- vegetation, productivity

Global change threats to wetlands

- Development
- Saltwater intrusion
- Temperature & CO₂

Definitions for today

Oxidation –

Reduction –

Reductant –

Oxidant –

Oxidation state (number) –

Oxidation-reduction potential –

Gibbs Free Energy –



Many elements in the periodic table can exist in more than one oxidation state. Oxidation states are indicated by Roman numerals (e.g. (+I), (-II), etc). The oxidation state represents the "electron content" of an element which can be expressed as the excess or deficiency of electrons relative to the elemental state.

If we want to determine whether a reaction is oxidation or reduction, we need to know the oxidation number of the element & how it changes

n.b., 'valence', 'oxidation number' and 'formal charge' are commonly used interchangeably, but are actually distinct from one another. For our purposes, we are only dealing with oxidation number **Rules for determining oxidation number of an element**

(1) Oxidation state of an element in its elementary state = 0
 e.g., Cl₂, Na, P...etc.

(2) Oxidation state of an element in a monatomic (one one atom) ion is equal to the charge on the ion
 e.g., Na⁺ = +1; Cl⁻ = -1; Fe³⁺ = +3

(3) Oxidation state of certain elements is the same in all, or almost all of their compounds

e.g., Group 1A elements: Li, Na, K, Rb, Cs =+1 Group 2A elements: Be, Mg, Ca, Sr, Ba, Ra = +2 Group VII b elements: F, CI, Br, I, At = -1 in binary compounds

Oxygen is almost always -2 (Except: when bonded to O or F) **H is almost always +1**; Except with a metal, e.g. NaH, CaH₂ is -1 (4) The sum of the oxidation states in a neutral species is = 0; In a charged ion it is equal to the charge on the ion *e.g.*, Na₂Se: Na = +1x2 = 2, thus Se = -2 MnO₄⁻: O = -2x4 = -8, thus Mn = 8-1 = 7

(5) Fractional oxidation numbers are possible. e.g., in Na₂S₄O₆ (sodium tetrathionate), S has an oxidation number of +10/4 O: 6(-2) = -12 Na: 2(+1) = 2 Residual = -10, which must be balanced by S: 4(+10/4) = +10

(6) The oxidation number is designated by:
 Arabic number below the atom, or
 Roman numeral or Arabic number after the atom (in parentheses)

Many elements in the periodic table can exist in more than one oxidation state. Oxidation states are indicated by Roman numerals (e.g. (+I), (-II), etc).

The oxidation state represents the "electron content" of an element which can be expressed as the excess or deficiency of electrons relative to the elemental state.

Oxidation States

Element	Oxidation State	<u>Species</u>
Nitrogen		NO_3^-
		NO_2^-
		N_2
		NH_3, NH_4^+
Sulfur		SO ₄ ²⁻
		$S_2O_3^{2-}$
		S°
		H_2S, HS^-, S^{2-}
Iron		Fe ³⁺
		Fe ²⁺
Manganese		MnO_4^{2-}
C		$MnO_{2}(s)$
		MnOOH (s)
		Mn ²⁺

Many elements in the periodic table can exist in more than one oxidation state. **Oxidation states** are indicated by Roman numerals (e.g. (+I), (-II), etc). The oxidation state represents the "**electron content**" of an element which can be expressed as the excess or deficiency of electrons relative to the elemental state.

O---- J-4- --- C4-4--

Oxidation States		
	Assign O to be (-I	I) & H to be (+I), calculate.
Element	Oxidation State	<u>Species</u>
Nitrogen	N (+V)	NO ₃ -
0	N (+III)	NO_2^-
	N (O)	N_2
	N (-III)	$\rm NH_3, \rm NH_4^+$
Sulfur	S (+VI)	SO ₄ ²⁻
	S (+II)	$S_2O_3^{2-}$
	S (O)	S°
	S(-II)	H_2S, HS^-, S^{2-}
Iron	Fe (+III)	Fe ³⁺
	Fe (+II)	Fe ²⁺
Manganese	Mn (+VI)	MnO_4^{2}
8	Mn (+IV)	$MnO_{2}(s)$
	Mn (+III)	MnOOH (s)
	Mn (+II)	Mn ²⁺

Balancing oxidation-reduction reactions

Conventionally always put the oxidised species on the left, the reduced species on the right

e.g. $MnO_{4^{-}(aq)} + CI_{(aq)} = Mn^{2+}_{(aq)} + CI_{2(g)}$

- 1. Separate the reaction into a reduction and oxidation part $MnO_{4^{-}(aq)} = Mn^{2+}_{(aq)}$ reduction $Cl_{(aq)} = + Cl_{2(g)}$ oxidation
- Balance each 1/2 reaction with respect to mass then with respect to charge. Use e⁻, H⁺, H₂O or OH⁻

$$2CI_{(aq)}^{-} = CI_{2(g)}^{-} mass$$

$$2CI_{(aq)}^{-} = CI_{2(g)}^{-} + 2e^{-} mass + charge$$

Balancing oxidation-reduction reactions

 $MnO_{4(aq)} = Mn^{2+}(aq) + 4H_2O$ (mass oxygen)

 $MnO_{4(aq)}^{-} + 8H^{+} = Mn^{2+}_{(aq)} + 4H_{2}O$ (mass oxygen+ hydrogen)

 $MnO_{4(aq)}^{-} + 8H^{+} + 5e^{-} = Mn^{2+}_{(aq)} + 4H_2O (mass + charge)$

3. Combine half reactions so electron gain equals loss

5*A = 10 e⁻; 2*B = 10 e⁻ i.e. 5*A +2*B

 $10 \text{ Cl}_{(aq)} + 2\text{MnO}_4 + 16\text{H}^+ = 5\text{Cl}_{2(g)} + 2\text{Mn}^{2+}_{(aq)} + 8\text{H}_2\text{O}$

4. Check for atom and charge balance

Oxic: $[O_2] > 100 \ \mu M$

- Oxidation of OM by molecular oxygen
- Relatively low OM (especially continental fluvial sediments) because the OM is partially oxidized prior to sediment burial
- Fe-oxy-hydroxide phases are not reduced -- they are transformed to *hematite* (Fe₂O₃) during diagenesis (*e.g.*, Red beds)
- Presence of oxidized Mn phases
- Usually light brown sediment

Suboxic: $100 \mu M > [O_2] > 1 \mu M$

Non-sulfidic environment -- $[H_2S] < 1 \mu M$

- Although the oxidation of OM consumes O₂, there is insufficient OM to generate much H₂S
- NO_3^- , MnO_2 , and FeOX reduction
- Fe²⁺ and Mn²⁺ increase and become supersaturated with respect to *siderite* (FeCO₃), *rhodochrosite* (MnCO₃), *glauconite* ((K,Na)(Fe³⁺,AI,Mg)₂(Si,AI)₄O₁₀(OH)₂), and *vivianite* (Fe₃(PO₄)₂·8(H₂O))
- Usually grey sediment

Anoxic: $[O_2] < 1 \mu M$

1) Sulfidic environments -- $[H_2S] > 1 \mu M$

Oxidation of OM is by sulfate reduction

- Greigite (Fe₃S₄) and mackinawite (Fe_{HX}S) are first formed metastably, then react with H₂S to produce iron monosulfide (FeS) and pyrite (FeS₂).
- Alabandite (MnS) is only stable under <u>very high</u> concentrations of H₂S -- thus *rhodochrosite* (MnCO₃) is more common
- Requires large OM deposition rate
- Usually black sediment

Redox Potential: The Fundamentals

- Redox potential expresses the tendency of an environment to receive or supply electrons
 - An oxic environment has high redox potential because
 O₂ is available as an electron acceptor

For example, Fe oxidizes to rust in the presence of O_2 because the iron shares its electrons with the O_2 :

$$4Fe + 3O_2 \rightarrow 2Fe_2O_3$$



 In contrast, an *anoxic* environment has low redox potential because of the absence of O₂

the more positive the potential, the greater the species' affinity for electrons and tendency to be reduced

A Simple Electrochemical Cell



- FeCl₂ at different Fe oxidation states in the two sides
- Wire with inert Pt at ends -voltmeter between electrodes
- Electrons flow along wire, and Cl⁻ diffuses through salt bridge to balance charge
- Voltmeter measures electron flow
- Charge remains neutral



 Container on right side is more oxidizing and draws electrons from left side

Electron flow and Cl⁻ diffusion continue until an equilibrium is established – steady voltage measured on voltmeter

If container on right also contains O_2 , Fe³⁺ will precipitate and greater voltage is measured

 $4Fe^{3+} + 3O_2 + 12e^{-}$ $\rightarrow 2Fe_2O_3(s)$

 The voltage is characteristic for any set of chemical conditions

Redox Potential in Nature

- A mixture of constituents, not really separate cells
- We insert an inert *Pt electrode* into an environment and measure the voltage relative to a standard electrode [Std. electrode = H₂ gas above solution of known pH (theoretical, not practical). More practical electrodes are calibrated using this H₂ electrode.]
 - Example: when O₂ is present, electrons migrate to the Pt electrode:

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O_2 + 4e^- + 4H^+ \rightarrow 2H_2O
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– The electrons are generated at the H₂ electrode:

 $2H_2 \rightarrow 4H^+ + 4e^-$

Voltage between electrodes measures the redox potential

Redox Potential in Nature



Schlesinger - Fig. 7.13

Redox Potential of a Reaction

• General reaction:

Oxidized species + e^- + $H^+ \leftrightarrow$ reduced species

Redox is expressed in units of "pe," analogous to pH:

 $pe = -log \{e^{-}\}$ (or Eh = 2.3 RT pE/F)

where [e-] is the electron concentration or activity

"pe" is derived from the equilibrium constant (K) for an oxidation-reduction reaction at equilibrium:

$$K = \frac{[reduced species]}{[oxidized species][e^-][H^+]}$$

Oxidized species + e^- + $H^+ \leftrightarrow$ reduced species

$$K = \frac{[reduced species]}{[oxidized species][e^-][H^+]}$$

$$\log K = \log [red] - \log [ox] - \log [e^{-}] - \log [H^{+}]$$
$$\log K = -p_{red} + p_{ox} + pe + pH$$

If we assume [oxidized] = [reduced] = 1 (*i.e.*, at standard state), then:

$$\log K = pe + pH$$

$\log K = pe + pH$

The <u>Nernst Equation</u> can be used to relate this equation to *measured Pt-electrode voltage (Eh, E_h, E_H)*:

$$pe \equiv pE = \frac{F}{2.3RT}Eh$$
 or Eh = 2.3 RT pE/F

where:

Eh = measured redox potential as voltage

R = the Universal Gas Constant (= 8.31 J K⁻¹ mol⁻¹)

T = temperature in degrees Kelvin

F = Faraday Constant (= 23.1 kcal V⁻¹ equiv⁻¹)

2.3 = conversion from natural to base-10 logarithms

- Reducing solution
 - Electron activity is high
 - pe is low
 - High tendency to donate electrons
- Oxidizing solution
 - Electron activity is low
 - pe is high
 - High tendency to accept electrons
- All very analogous to treatment of pH

Eh- pH (pe – pH) Diagrams

- Used to show equilibrium speciation for reactants as functions of Eh and pH
- Red lines are practical Eh-pH limits
 on Earth

•
$$pE = pe = \frac{F}{2.3RT}Eh$$



Eh-pH diagram for H_20

Eh-pH diagrams describe the thermodynamic stability of chemical species under different biogeochemical conditions

1.2 Fe⁺³ aq 1.0 FeOH⁺² aq **O**₂ Fe(OH)₂⁺ aq H_2O dE/dpH = 0.8 --0.059 Fe(OH)₃ 0.6 -(volts) Example – predicted 0.4 · Fe⁺² aq stable forms of Fe in 0.2 aqueous solution: 0.0 H_2O -0.2 -Η2 Fe₃(OH)₈ -0.4 -Fe(OH)₂ -0.6 Diagram is for 25 degrees C

2

3

9

8

pН

10

11

12

Example –

Oxidation of H₂S released from anoxic sediments into oxic surface water:



Redox Reactions in Nature

Example: net reaction for aerobic oxidation of organic matter:

$$CH_2O + O_2 \rightarrow CO_2 + H_2O$$

In this case, oxygen is the *electron acceptor* – the half-reaction is:

$$O_2 + 4H^+ + 4e^- \rightarrow 2H_2O$$

- Different organisms use different electron acceptors, depending on availability due to local redox potential
- The more oxidizing the environment, the higher the energy yield of the OM oxidation (the more negative is ΔG, the Gibbs free energy)

Free Energy and Electropotential

- Talked about electropotential (aka emf, Eh) --> driving force for e⁻ transfer
- How does this relate to driving force for any reaction defined by ∆G_r??

∆G_r = - nℑE

Where n is the # of e-'s in the rxn, ℑ is Faraday's constant (23.06 cal V⁻¹), and E is electropotential (V)

 pe for an electron transfer between a redox couple analogous to pK between conjugate acidbase pair

- The higher the energy yield, the greater the benefit to organisms that harvest the energy
- In general:
 - There is a temporal and spatial sequence of energy harvest during organic matter oxidation
 - Sequence is from the use of high-yield electron acceptors to the use of low-yield electron acceptors

The greater the difference in pe between the oxidizing & reducing agents, the greater the free energy yield for the reaction

Sets up a sequence of favorable oxidants for organic matter oxidation

Organic matter oxidation by O_2 is greatest free energy yield

Why is organic matter such a good electron donor?



Z-scheme for photosynthetic electron transport



energy from sun converted to C-C, energy rich, chemical bonds

Environmentally Important Organic Matter Oxidation Reactions

Reducing Half-reaction	E _h (V)	ΔG
Reduction of O ₂		
$O_2 + 4H^+ + 4e^ > 2H_2O$	+0.812	-29.9
Reduction of NO ₃ ⁻		
$2NO_3^- + 6H^+ + 6e^> N_2 + 3H_2O$	+0.747	-28.4
Reduction of Mn (IV)		
$MnO_2 + 4H^+ + 2e^> Mn^{2+} + 2H_2O$	+0.526	-23.3
Reduction of Fe (III)		
$Fe(OH)_3 + 3H^+ + e^> Fe^{2+} + 3H_2O$	-0.047	-10.1
Reduction of SO ₄ ²⁻		
SO ₄ ²⁻ + 10H ⁺ + 8e ⁻ > H ₂ S + 4H ₂ O	-0.221	-5.9
Reduction of CO ₂		
$CO_2 + 8H^+ + 8e^> CH_4 + 2H_2O$	-0.244	-5.6

Π C REASING Π Z RGY YIELD