

CONSOLIDATION MECHANICS REVIEW (41)

A Review of key points

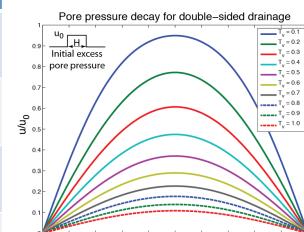
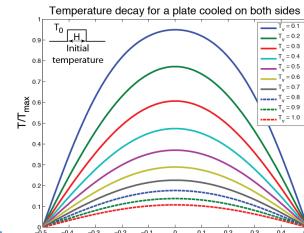
B Worked examples

II Review of key points

A Heat flow analogous to
flow of fluid during
consolidation

$$\frac{\partial T}{\partial t} = \alpha \frac{\partial^2 T}{\partial z^2} \quad \frac{\partial u_{excess}}{\partial t} = C_v \frac{\partial^2 u_{excess}}{\partial z^2}$$

Heat flow	Fluid flow
T Temperature	u_{excess} Excess pore pressure
t Time	t Time
α Thermal diffusivity	C_v Coefficient of consolidation
z position	z position



Here $T_v = t^* = t/(H^2/C_v)$

II Review of key points

B Darcy's law $q = \frac{Q}{A} = -k \frac{\partial H}{\partial x}$
 where

q = flux (m/sec)

Q = discharge rate (m^3/sec)

A = area (m^2)

K = hydraulic conductivity (m/sec)

$\partial H/\partial x$ = head gradient (dimensionless)

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II Review of key points

C Darcy's law $Q = -kA \frac{\Delta H}{\Delta L}$
 (alternative form)

where

Q = discharge rate (m^3/sec)

k = hydraulic conductivity (m/sec)

A = area (m^2)

ΔH = change in head (m)

ΔL = length of flow path (m)

k is a function of the fluid and the porous medium

Relationship between hydraulic conductivity and intrinsic permeability

$$k = K \frac{\rho g}{\mu}$$

$$K = k \frac{\mu}{\rho g}$$

where

k = hydraulic conductivity (m/sec)

K = intrinsic permeability (m^2)

ρ = density of the fluid (kg/m^3)

g = acceleration due to gravity (m/sec^2)

μ = dynamic viscosity of fluid [$\text{kg}/(\text{m}\cdot\text{sec})$]

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II Review of key points

- B Controls on 1-D consolidation time (t_c) of soils and sediments
- 1 The time for consolidation should scale with the volume of water squeezed out of the soil
 - 2 The volume of expelled water is the product of three terms
 - a Effective stress change ($\Delta\sigma'$)
 - B Compressibility of the solid skeleton (i.e., m_v , the coefficient of volume change)
 - C Volume of the consolidating layer (or height H of the layer)

$$t_c \propto \Delta\sigma' m_v H$$

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II Review of key points

- B Controls on 1-D consolidation time (t_c) of soils and sediments (cont.)
- 3 The time [for consolidation] should be inversely proportional to how fast the water can flow through the consolidating layer
 - 4 From Darcy's law: the velocity of flow is scales with the product of the permeability (k) and the hydraulic gradient
 - 5 The hydraulic gradient matches the excess fluid pressure lost within the layer divided by the distance the pore fluid flows (H)
 - 6 The excess fluid pressure loss scales with $\Delta\sigma'$

$$t_c \propto \frac{1}{k\Delta\sigma'/H} = \frac{H}{k\Delta\sigma'}$$

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II Review of key points

B Controls on 1-D consolidation time (t_c) of soils and sediments (cont.)

- 7 t_c increases with increasing compressibility (i.e., with increasing m_v)
- 8 t_c Increases rapidly with increasing volume (height) of the soil/sediment mass (H)
- 9 t_c Decreases with increasing permeability (k)
- 10 t_c is *independent* of the magnitude of the effective stress change, assuming m_v is a constant

$$t_c \propto \frac{\Delta\sigma' m_v H}{k\Delta\sigma'/H} = \frac{m_v H^2}{k}$$

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II Review of key points

C Dimensional analysis illuminates consolidation time

- 1 1-D consolidation equation

$$\frac{\partial u_e}{\partial t} = C_v \frac{\partial^2 u_e}{\partial x^2}$$

$$\frac{\text{pressure}}{\text{time}} = [C_v] \frac{\text{pressure}}{\text{length}^2} \Rightarrow [C_v] = \frac{\text{length}^2}{\text{time}}$$

where

 C_v = coefficient of consolidation

$$= \frac{k}{\rho_{\text{water}} g m_v}$$

 m_v = coefficient of volume change

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II Review of key points

2 The dimensionless time used in dimensionless expressions for consolidation is found in the same way as in lecture 38

$$\frac{\partial T}{\partial t} = \alpha \frac{\partial^2 T}{\partial x^2}$$

$$t_c \approx \frac{H^2}{\alpha}$$

$$t^* = \frac{t}{t_c} = \frac{t}{\left(\frac{H^2}{\alpha}\right)}$$

$$\frac{\partial u_e}{\partial t} = C_v \frac{\partial^2 u_e}{\partial x^2}$$

$$t_c \approx \frac{H^2}{C_v}$$

$$t^* = \frac{t}{t_c} = \frac{t}{\left(\frac{H^2}{C_v}\right)}$$

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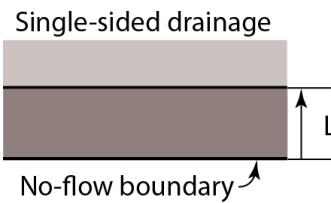
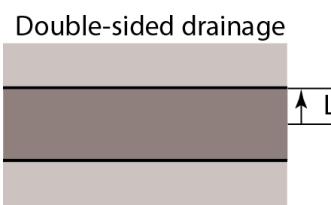
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II Review of key points

D The key length scale (L) is the flow path length

- 1 For double-sided drainage, L = half the layer thickness
- 2 For single-sided drainage, L = the whole layer thickness



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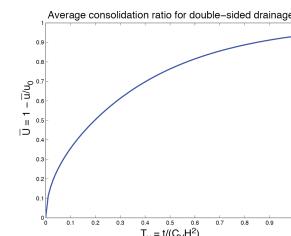
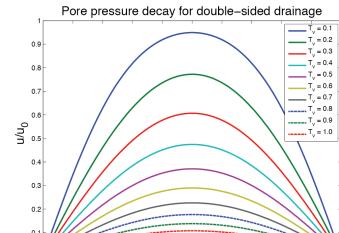
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III Calculating consolidation

$$U(t) = \frac{\Delta H(t)}{\Delta H(t \rightarrow \infty)} = \frac{\Delta u_{excess}}{\Delta u_{excess(max)}} = 1 - \frac{u_{excess}}{u_0}$$

E Consolidation ratio for double-sided drainage of a uniform excess pressure pulse* [for analytical solution see Terzaghi (1943) or Lambe and Whitman (1969)]

- 1 For $T_v = t/(C_v H^2) > 0.1$
 - a Excess pore pressure distribution (u) is approximately parabolic
 - b \bar{U} (ave. consolidation ratio)
 - i $\bar{U} = 1 - \frac{\bar{u}}{u_0}$
 - ii $\bar{U} = 1 - \frac{\text{area beneath } u\text{-curve}}{\text{area beneath boxcar function}}$
 - iii $\bar{U} \approx 1 - \frac{2}{3} \frac{u_{max}}{u_0}$
 - 2 ~92% consolidation at $T_v=1$



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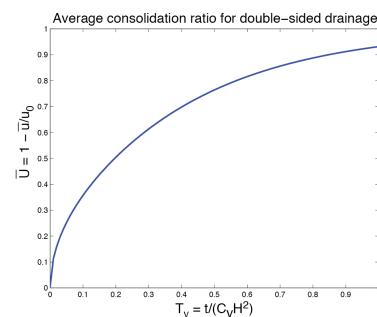
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II Review of key points

E Average consolidation ratio (\bar{U})

- 1 $\bar{U}(t^*=1) \approx 92\%$
- 2 $\bar{U}(t^*=3) \approx 99\%$
- 3 About 92% of the ultimate primary consolidation occurs by the dimensionless time $t^* = 1$



$$t^* = \frac{t}{\left(\frac{x_{max}^2}{C_v} \right)}$$

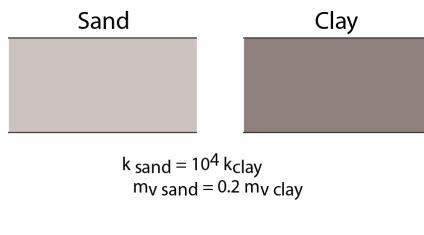
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III Examples

A Consider layers of clay and sand, each 10' thick, and suppose that the coefficient of volume change ("compressibility") of the sand is 1/5 that of the clay, and the permeability of the sand is 10,000 times that of the clay.



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III Examples

- 1 What is the ratio of the consolidation times of the sand and clay?

$$\frac{t_{\text{clay}}}{t_{\text{sand}}} = \frac{H_{\text{clay}}^2 m_v(\text{clay})/k_{\text{clay}}}{H_{\text{sand}}^2 m_v(\text{sand})/k_{\text{sand}}} = \frac{m_v(\text{clay})/k_{\text{clay}}}{m_v(\text{sand})/k_{\text{sand}}} = \frac{5/1}{1/10,000} = 50,000$$

- 2 If a 10'-thick layer of clay reaches 90% consolidation in 10 years, how long would it take for a clay layer 40' thick to reach that level of consolidation?

$$\frac{t_{40'}}{t_{10'}} = \frac{m_{v40'} H_{40'}^2 / k_{40'}}{m_{v10'} H_{10'}^2 / k_{10'}} = \frac{H_{40'}^2}{H_{10'}^2} = \frac{40^2}{10^2} = 16 \quad 16 \times 10 \text{ years} = 160 \text{ years}$$

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III Examples

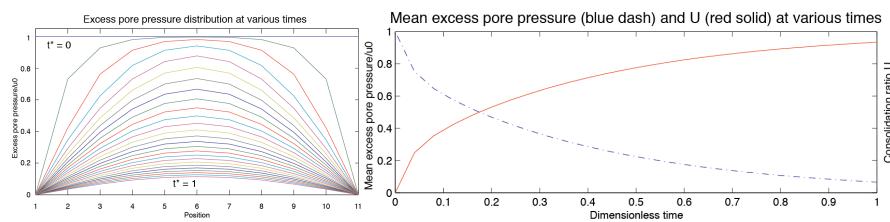
- B How long will a layer of clay take to reach 90% consolidation if the initial excess pore pressure distribution is constant across the layer and the layer is drained from its top and bottom?
One dimensionless time unit (t^*) is given by = 1.

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C Calculating consolidation for double-sided drainage 4 Matlab script transient_heat4.m



So after one dimensionless time step [i.e. $t^* = t/(H^2/C_v) = 1$], about 92% (i.e., roughly 90%) of the ultimate consolidation will have occurred

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C Calculating consolidation for double-sided drainage

3 Matlab function consolidation_v2.m

```

function [U,u] = consolidation_v2(t,Cv,H)
% function [U,u] = consolidation_v2(t,Cv,H)
% Calculates time factor for 1-D consolidation for a uniform excess
% pressure pulse of unit magnitude (i.e., a boxcar function)
% Output arguments
% U = consolidation ratio
% u = excess pore pressure
% Input arguments
% t = time
% Cv = coefficient of consolidation
% H = layer thickness
% Example
% t=0.01:1; Cv = 1; H = 1;
% [U,u] = consolidation_v2(t,Cv,H);
% Calculate time scale Tv
Tv = Cv^2/(H^2);
z = 0.01:H:2*H;
u = zeros(length(z),length(t));
% Calculate pore pressure (u) by truncated Fourier series (101 terms)
% using eq. (3a) of Terzaghi (1943, p. 274)
for j = 1:length(t);
    for i = 1:length(z);
        for k = 1:101
            N = k-1;
            n = 2*N+1;
            uu(i,j)=(4/pi)*(1/n)*sin(n*pi*(z(i)/(2*H)))*exp(-Tv(j)*(n*pi/2).^2);
        end
        u(i,j)=sum(uu);
    end
    % Calculate consolidation ratio U by trapezoidal integration
    U(j) = 1 - sum(u(:,j))/(length(z)-1);
end

```

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C Calculating consolidation for double-sided drainage

4 Matlab script transient_heat4.m

```

% Matlab script transient_heat4.m
% Solves for 1-D transient heat flow by finite differences
% H = length of longest flow path
a = 11; % # of rows in the T matrix (# of points in space)
b = 26; % # of columns in the T matrix (# of points in time)
numit = 21; % # of iterations at each time step
T0 = 0; % Temperature at one end of rod
T1 = 1; % Temperature at other end of rod

% Initialize the "Temperature matrix"
T=zeros(a,b); % First index is position, second is time step
% Set the initial temperature distribution
T(:,1) = (linspace(T1,T0,a))' % Constant initial temperature

% Solve by finite difference method
for j=2:b; % j is the index for time step
    for k = 1:numit;
        for i = 2:a-1; % i is index for position
            T(i,j) = 0.25*( T(i+1,j-1)+T(i-1,j-1)+T(i+1,j)+T(i-1,j) );
        end
    end
end

```

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References

- Lambe, T.W., and Whitman, R.V., 1969, Soil Mechanics: Wiley, New York, 553 p.