Near-Axis Seamount Volcanism







Controls on Seamount Formation Overall abundance controlled by variations in regional magmatic budget at both slow- [Jaroslow, 2000] and fast-spreading ridges [Gomez, 2000] Preferentially located near bends or discontinuities in ridge crest where crustal vulnerability is relatively high, and local topographic highs along ridge where local magma supply is high Do not appear to be controlled by lithospheric cracking or thermal contraction [Forsyth, 2005], but may be localized by fracture zones or other linear discontinuities [Scheirer, 1995] Seamount chains may be created by moving asthenospheric melt anomalies [Forsyth, 2005]; isolated volcanoes likely related to axial upwelling [White, 1998] Result from shallower melting than axial lavas, potentially from near the cooler edges of the upwelling zone [Batiza, 1990] Inferred width of zone where volcanoes initially form varies with spreading rate: <4 km for MAR, <20 km for NEPR, <30 km for SEPR, primarily on lithosphere <0.2 Ma age (<4-6 km brittle thickness), regardless of spreading rate [White, 1998] Bear Seamount Clipperton FZ/ Lamont Seamounts NW Atlantic











	MOR Seamounts: Intermediate Spreading
•	Not nearly as much info as EPR seamounts, mostly focus on large chain-
	forming volcanoes along Juan de Fuca [Clague et al., 2000] and Gorda
	Ridges [Davis et al., 2000]
•	Volumes range from 11 to 187 km ³ , diameter from ~3-10 km, and height
	from ~500-2000 m for the surveyed chains
•	President Jackson seamount chain (Gorda):
	 Lava compositions range from highly depleted to moderately enriched MORB
	 Higher MgO than axial lavas
	 Phenocrysts in equilibrium with melt and rare zoning indicate short residence time in crustal magma chambers
	 Often have multiple nested calderas and pit craters, implying episodic melt supply on time scale of 1000-10,000 yrs between eruptions
•	Vance Seamounts (Juan de Fuca):
	 Some edifices are cut by ridge-parallel faults, implying formation within ~30-40 km of axis
	 Faults are interpreted as the pathway for seamount lavas to the seafloor
	 Seamounts must have magma chambers because calderas exist, but melts must pass through quickly to retain primitive character
	 Magma chamber volumes similar to HI volcanoes: ~3-5 km diameter, ~3 km³ volume, 0.8 to 3 km below the seafloor
	 Chains are oblique to spreading direction and plate motion, align with asthenospheric flow
	 Lifespan of volcanoes is ~75 to 95 kyr (much shorter than EPR, based on asthenospheric flow rates, seems too short to construct a large edifice)





Summary
 Near-axis seamounts bypass the homogenizing processes within the axial magma chamber, allowing us to sample and understand small- scale heterogeneity within the mantle underneath oceanic spreading systems
 Form within a few km of the axis, continue growing to 10's of km off- axis at intermediate and fast-spreading ridges, only grow within the inner rift valley at slow ridges
 Lavas are more primitive and higher MgO in general, vary between depleted NMORB and enriched EMORB, more variable within and between seamounts than axial lavas
 Seamount production correlates with changes in melt supply at different scales:
 Spreading rate-related changes in melt supply at the regional to global scale for MOR's, proximity to volcanic arc controls melt supply in backarc
 Local changes in melt supply at the first and second-order segment scale: increased toward the center, decreased toward ends, local increases near discontinuities or fracture zones
 Near-axis seamounts in the backarc setting seem to be formed by a similar mechanism to MOR's

Areas in need of further study Range of eruptive styles, explosive vs. effusive, both? Volume of individual eruptions Timing/frequency of eruptions More age data in general Hydrothermal systems/biota More seismic data to find extent of low-velocity zone under all types of ridges and its relationship to seamounts