

1 Relationship of environmental relative humidity with North 2 Atlantic tropical cyclone intensity and intensification rate

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5 Received 13 August 2012; revised 21 September 2012; accepted 23 September 2012; published XX Month 2012.

6 [1] Quantifying the relationship of large-scale environ-
7 mental conditions such as relative humidity with hurricane
8 intensity and intensity change is important for statistical
9 hurricane intensity forecasts. Our composite analysis of 9
10 years of Atmospheric Infrared Sounder (AIRS) humidity data
11 spanning 198 Atlantic tropical cyclones (TCs) shows that
12 environmental relative humidity (ERH) above the boundary
13 layer generally decreases with time as TCs evolve. Near the
14 surface, ERH stays approximately constant. ERH generally
15 increases with increasing TC intensity and intensification
16 rate. Rapidly intensifying TCs are associated with free tro-
17 pospheric ERH more than 10% (relative to the averaged ERH
18 for all TCs) larger than that for weakening TCs. Substantial
19 azimuthal asymmetry in ERH is also found, especially for the
20 TCs attaining the highest intensities and largest intensifica-
21 tion rates at distances greater than 400 km away from the TC
22 center. In the front-right quadrant relative to TC motion,
23 rapid intensification is associated with a sharp gradient of
24 ERH in the upper troposphere, with a decrease from the near
25 to the far environment between 400 hPa and 300 hPa. The
26 ERH gradient weakens with the decrease of intensification
27 rate. This radial ERH gradient might be a useful predictor for
28 the statistical forecast of TC intensification. **Citation:** Wu,
29 L., H. Su, R. G. Fovell, B. Wang, J. T. Shen, B. H. Kahn, S. M.
30 Hristova-Veleva, B. H. Lambriksen, E. J. Fetzer, and J. H. Jiang
31 (2012), Relationship of environmental relative humidity with
32 North Atlantic tropical cyclone intensity and intensification rate,
33 *Geophys. Res. Lett.*, 39, LXXXXX, doi:10.1029/2012GL053546.

34 1. Introduction

35 [2] While our understanding of tropical cyclones (TCs) has
36 improved tremendously in the past several decades, forecasts
37 of TC genesis, spin-up and subsequent (especially sudden)
38 intensity changes still present significant challenges. Official
39 intensity forecasts from the National Hurricane Center
40 (NHC) for Atlantic and Eastern North Pacific TCs have not
41 shown much improvement in the last 20 years [*DeMaria*

et al., 2007]. This is because TCs are sensitive to many fac- 50
tors, within the storm and in its surrounding environment. 51
For example, TC structure and intensity are sensitive to ver- 52
tical wind shear in the environment [*DeMaria*, 1996; *Frank* 53
and Ritchie, 2001; *Zehr*, 2003], which may be poorly fore- 54
casted by operational and research models. Also, relatively 55
subtle variations in sea-surface temperature (SST) or ocean 56
heat content can cause a TC intensity to shift several cate- 57
gories on the Saffir-Simpson scale within a short period of 58
time [*Sun et al.*, 2007]. 59

[3] The available moisture of the TC's environment 60
represents another poorly understood influence on intensity, 61
thereby presenting a limit to predictability. While high mid- 62
tropospheric relative humidity (RH) appears to be necessary 63
for rapid intensification and the attainment of maximum 64
intensity [e.g., *Kaplan and DeMaria*, 2003; *Emanuel et al.*, 65
2004; *Hendricks et al.*, 2010; *Kaplan et al.*, 2010], dry air 66
intrusions have a negative influence on TC intensification as 67
dry air ingestion promotes the formation of cold downdrafts, 68
which transport low θ_e air into the sub-cloud layer and storm 69
inflow [e.g., *Emanuel*, 1989]. In an idealized modeling study, 70
Braun et al. [2012] showed that low humidity air reaching 71
the inner core induces asymmetric convective activity which 72
weakens TCs [e.g., *Nolan and Grasso*, 2003; *Nolan et al.*, 73
2007]. They further showed that the time for TCs to reach 74
maturity varies with the proximity of dry air to the center of 75
circulation. When dry air is located 270 km away or further 76
from the center of the vortex, its impact on TC intensity is 77
insignificant. 78

[4] Some studies [e.g., *Barnes et al.*, 1983; *Wang*, 2009] 79
have shown that substantial and extensive moisture may also 80
promote a net negative influence on TC strength by facilitat- 81
ing the formation of TC rainbands. The idealized modeling 82
study of *Hill and Lackmann* [2009], which varied the envi- 83
ronmental RH (ERH) in the region ≥ 100 km beyond the TC 84
core, suggests that larger ERH results in the establishment of 85
wider TCs with more prominent outer rainbands. However, 86
TC development, as measured by time series of maximum 87
10 m wind speeds, was nearly insensitive to ERH despite the 88
variation in rainband activity. 89

[5] *Kaplan and DeMaria* [2003] examined the mid- 90
tropospheric (850–700 hPa) ERH relation with rapidly 91
intensifying (RI) TCs in the North Atlantic basin using the 92
NHC HURDAT file [*Jarvinen et al.*, 1984] and the Statistical 93
Hurricane Intensity Prediction Scheme (SHIPS) [*DeMaria* 94
and Kaplan, 1999] database. *Hendricks et al.* [2010] con- 95
ducted composite analyses using the Navy Operational 96
Global Atmospheric Prediction System (NOGAPS) global 97
analysis. Both studies found that RI events over the Atlantic 98
basin are associated with larger RH in the middle troposphere 99
than non-RI events. 100

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