

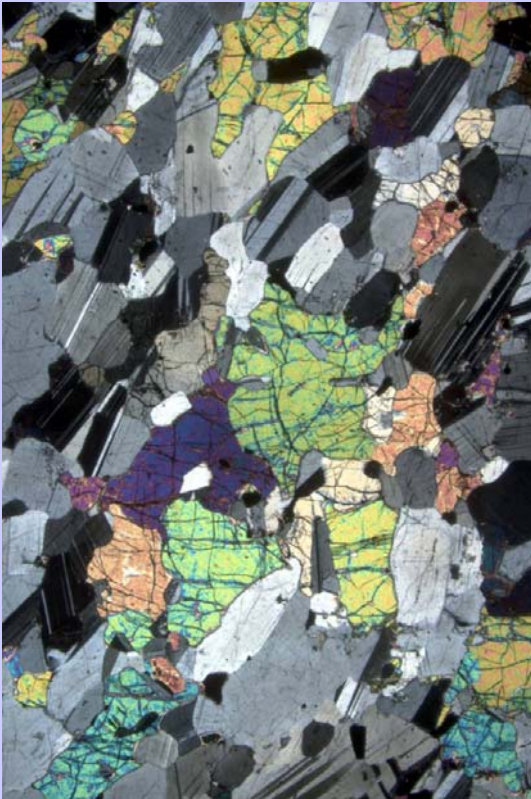


# Optical Mineralogy in a Nutshell

## Use of the petrographic microscope



# Why use the petrographic microscope?



- Identify minerals (*no guessing!*)
- Determine rock type
- Determine crystallization sequence
- Document deformation history
- Observe frozen-in reactions
- Constrain P-T history
- Note weathering/alteration
- Fun, powerful, and cheap!

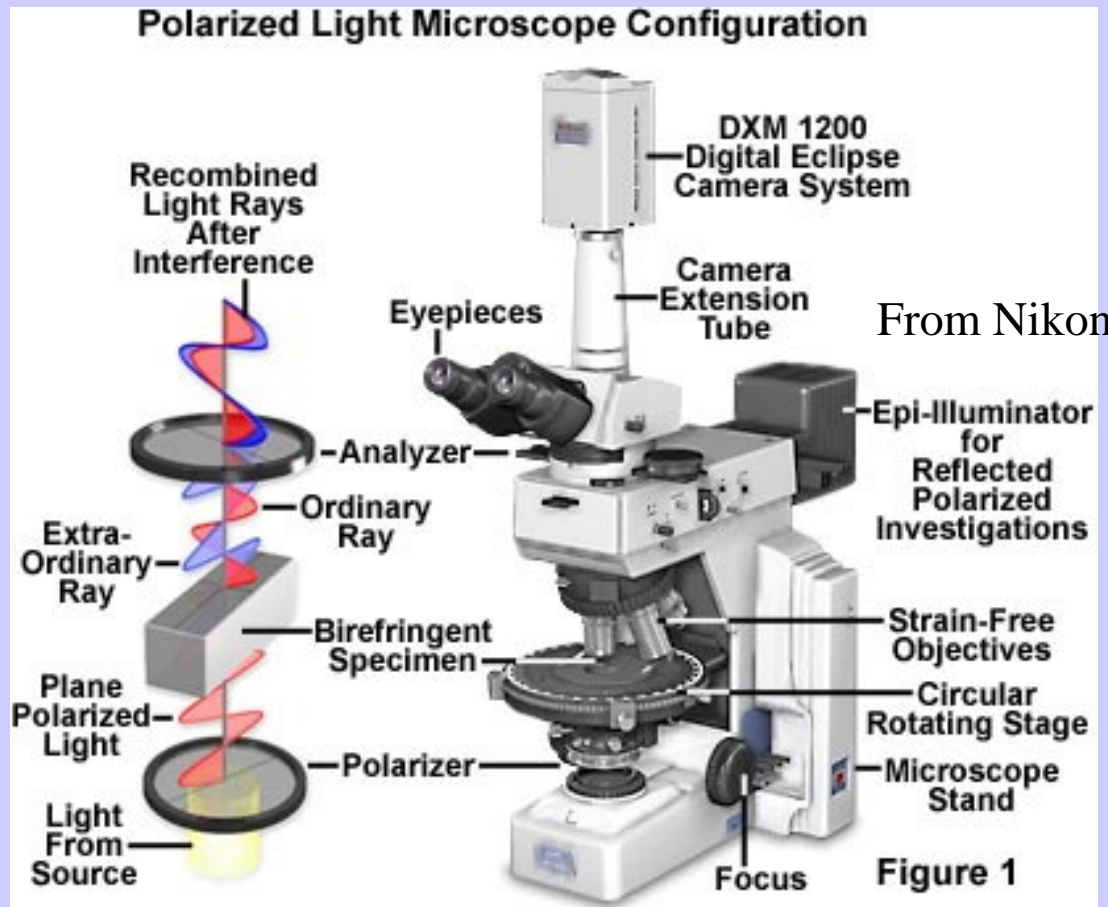
# The petrographic microscope



Also called a  
polarizing  
microscope

In order to use the scope, we need to understand a little about the **physics of light**, and then learn some **tools and tricks**...

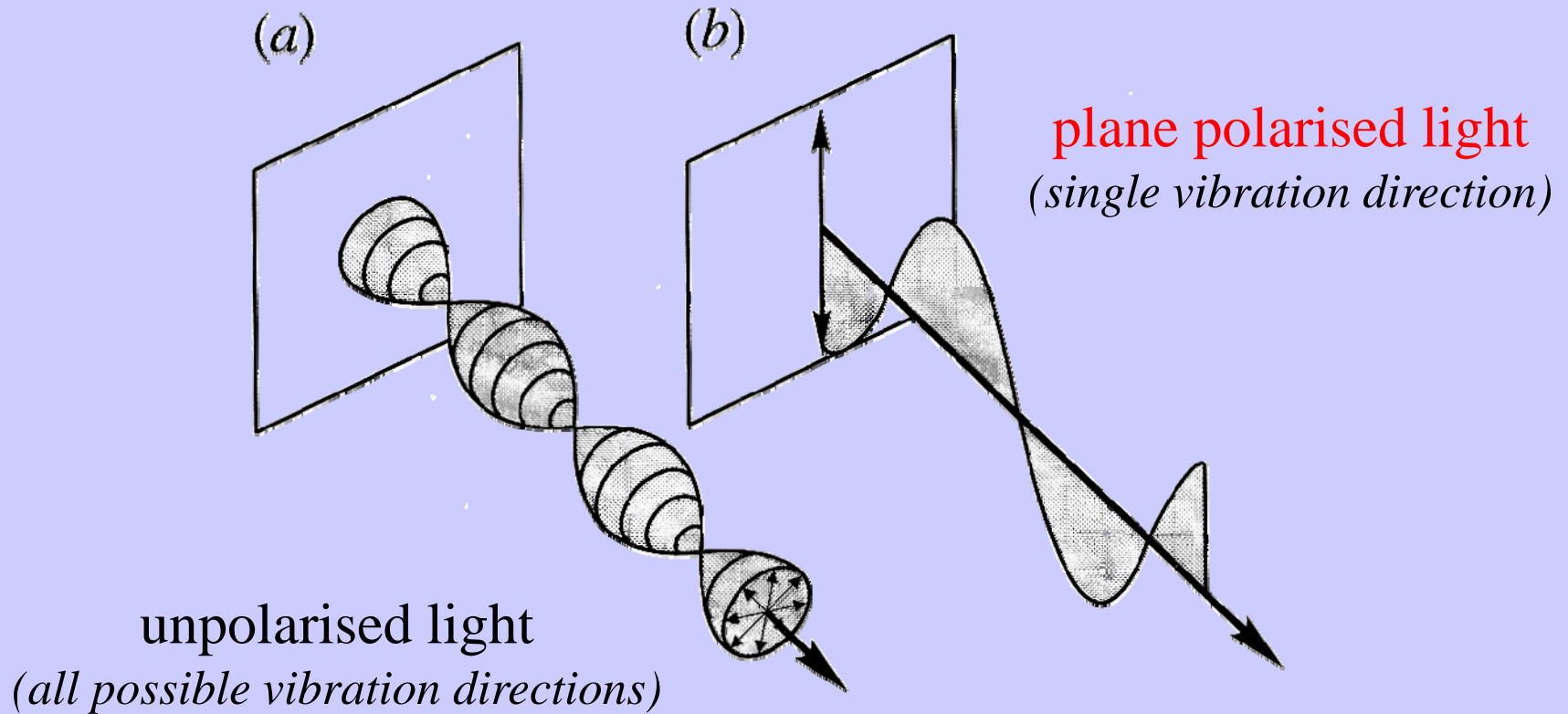
# Polarized Light Microscopy



Isotropic materials, which include gases, liquids, unstressed glasses and cubic crystals, demonstrate the same optical properties in all directions. They have only one refractive index and no restriction on the vibration direction of light passing through them.

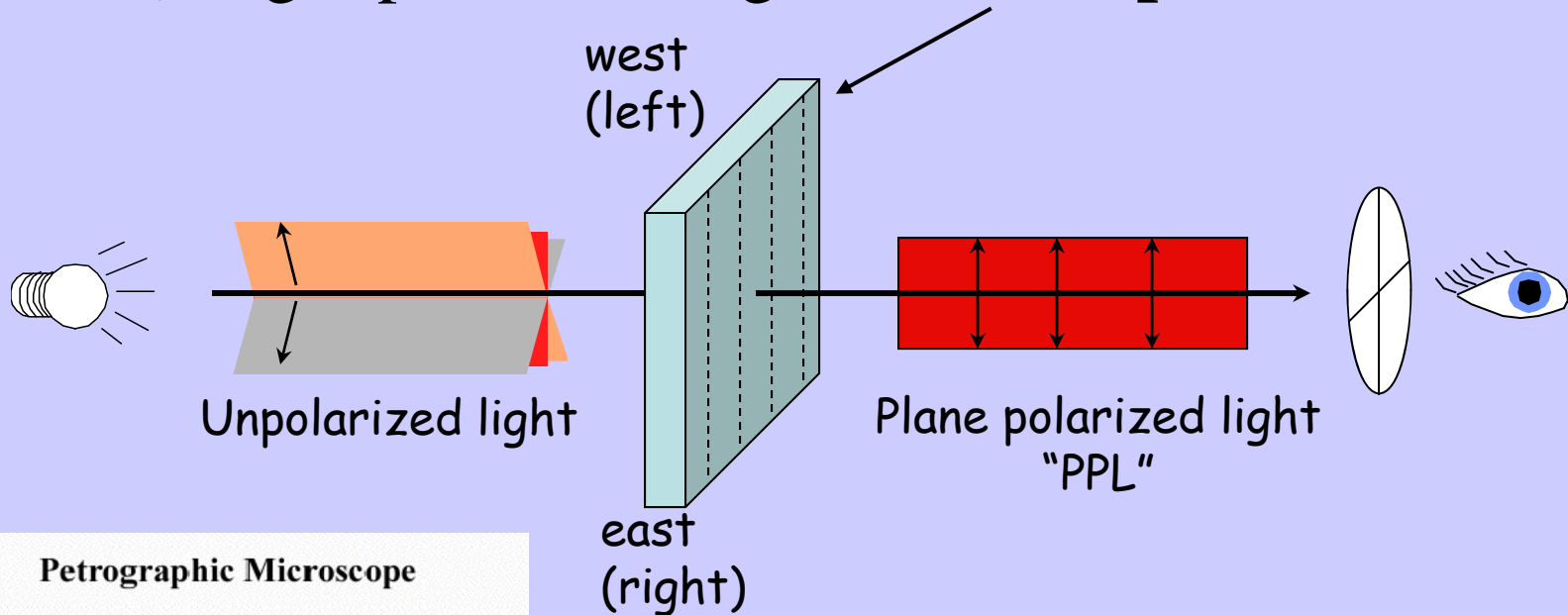
Anisotropic materials, in contrast, which include 90 percent of all solid substances, have optical properties that vary with the orientation of incident light with the crystallographic axes. Anisotropic materials act as beam splitters and divide light rays into two parts. The technique of polarizing microscopy exploits the interference of the split light rays, as they are re-united along the same optical path to extract information about these materials.

# What happens as light moves through the scope?



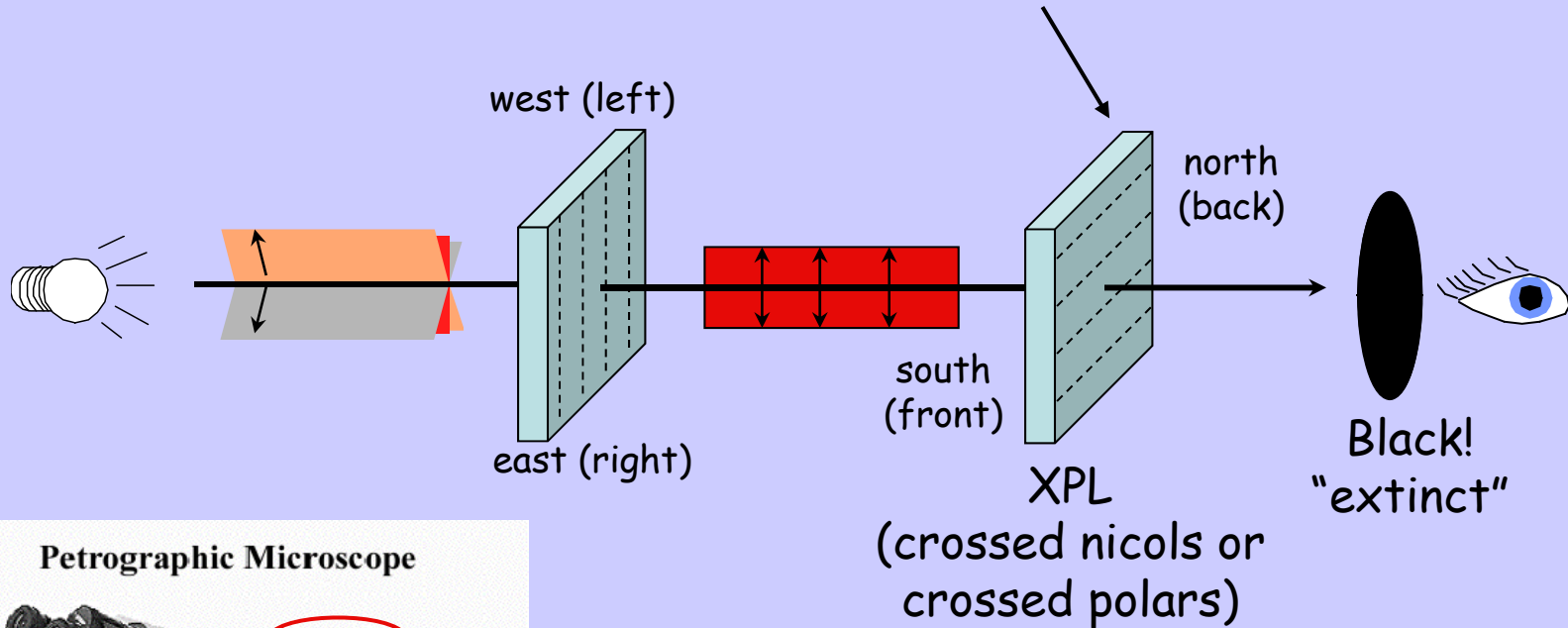


# 1) Light passes through the **lower polarizer**



Only the component of light vibrating in E-W direction can pass through lower polarizer -  
**light intensity decreases**

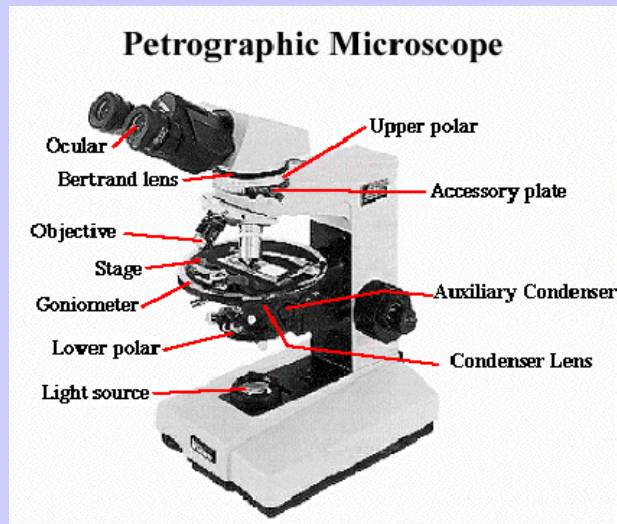
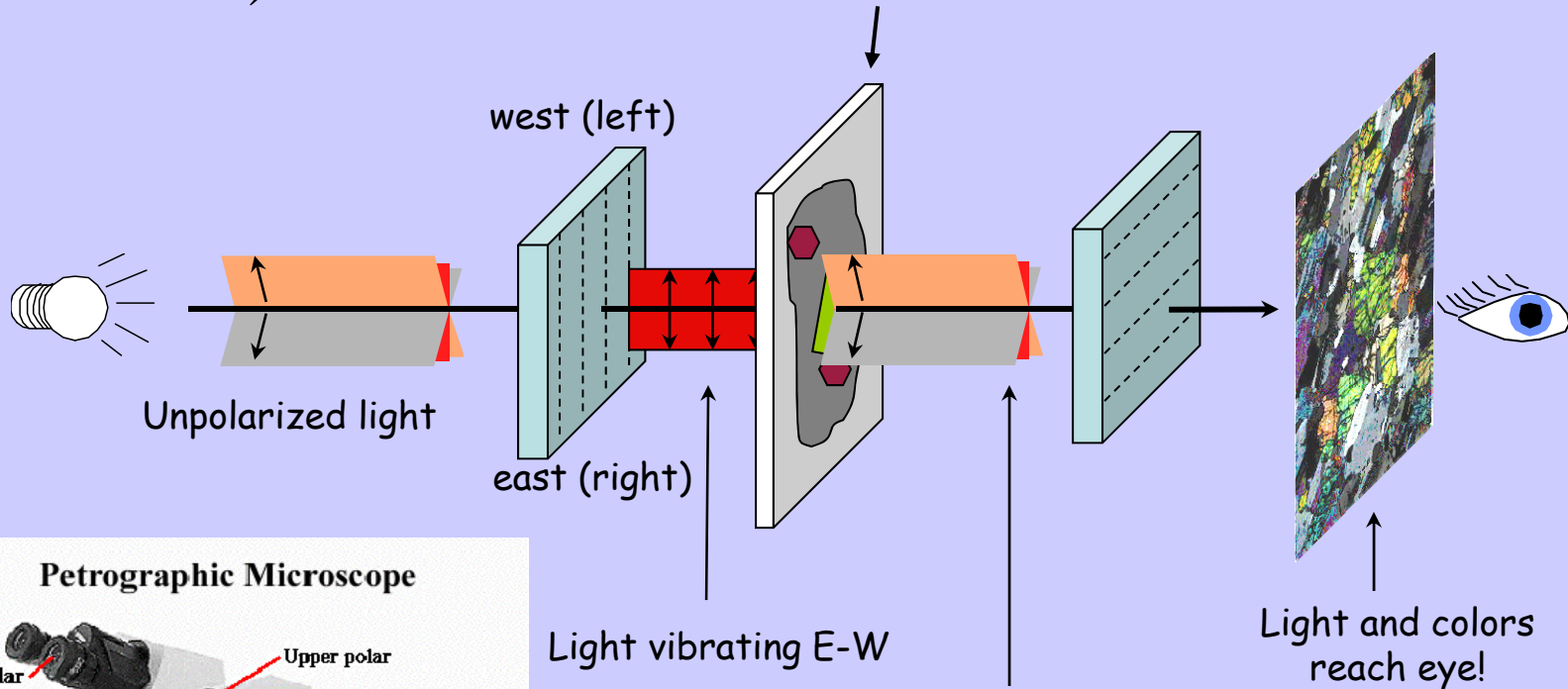
## 2) Insert the upper polarizer



Now what happens?  
What reaches your eye?

Why would anyone design a microscope that prevents light from reaching your eye?

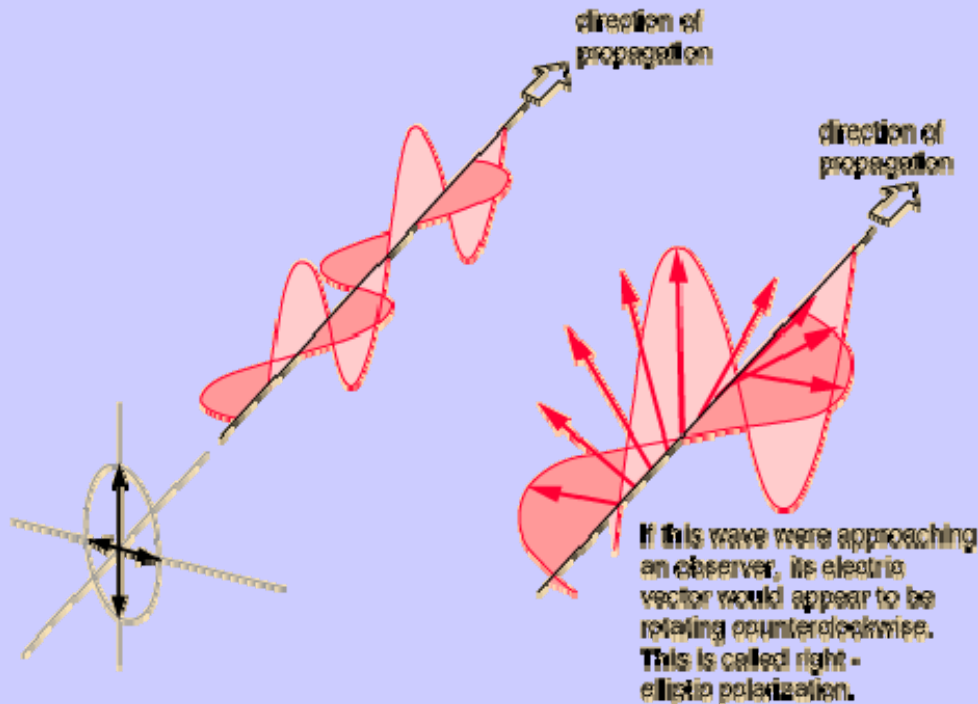
### 3) Now insert a **thin section** of a rock



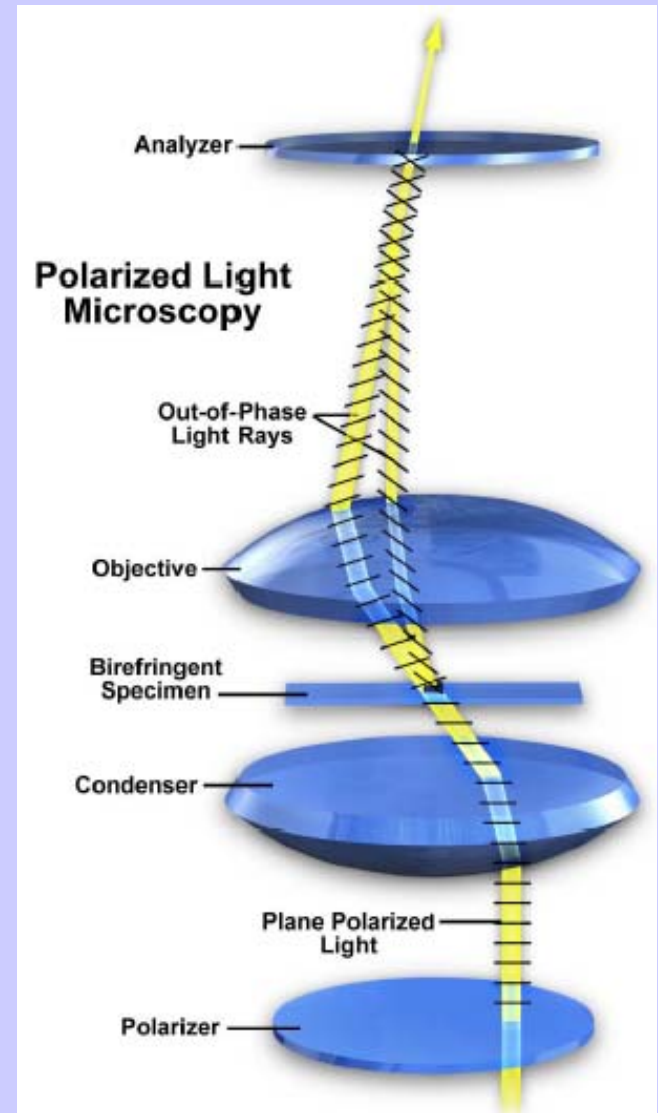
How does this work?



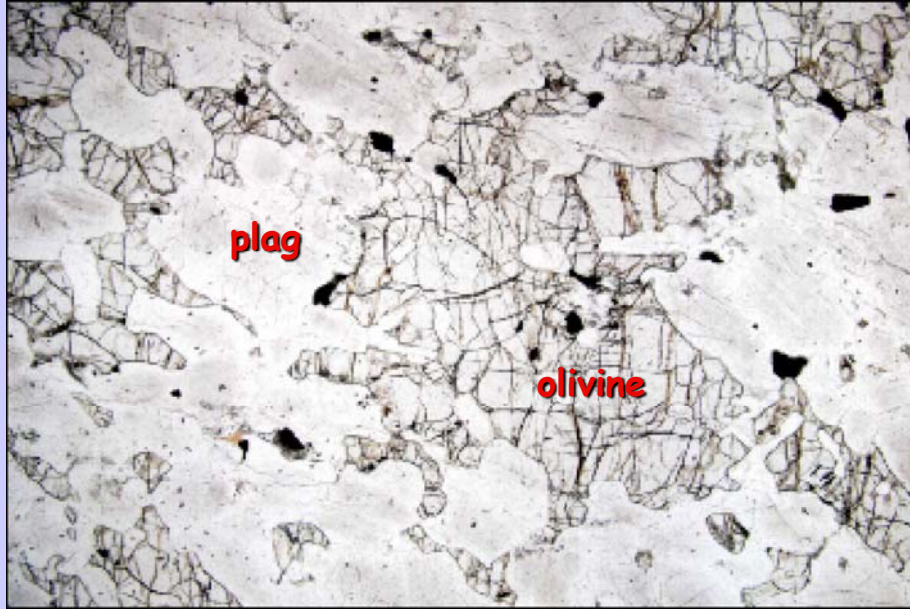
# Polarized Light Microscopy



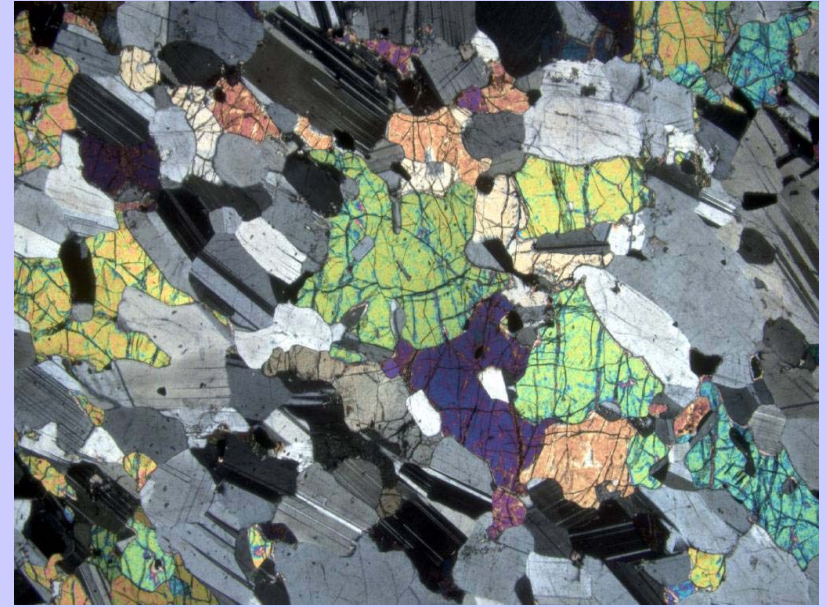
- The Light source is polarized before entering the specimen
- The analyzer pass the light with polarization angle perpendicular to the source light
- The contrast of birefringent material in the specimen would be enhanced



Conclusion has to be that minerals somehow **reorient** the planes in which light is vibrating; some light passes through the upper polarizer



PPL



XPL

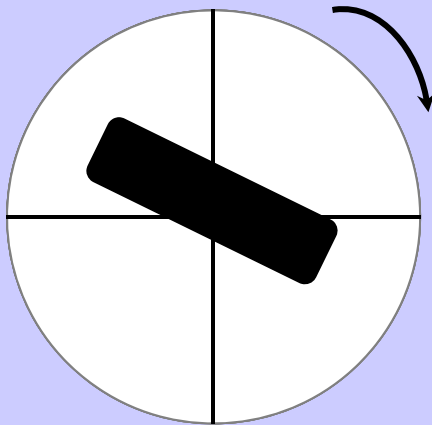
**Minerals act  
as magicians!**



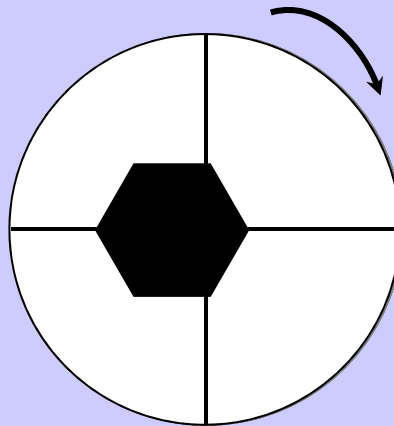
But, note that some minerals are better magicians than others (i.e., some grains stay dark and thus can't be reorienting light)

# Note the effect of **rotating the stage**

Most mineral grains change to black as the stage is rotated; they go black 4 times in a 360° rotation - **exactly every 90°**



Such minerals are **anisotropic**



However, glass, a few isotropic minerals, liquids (e.g., fluid inclusions), and gases (e.g., air bubbles) stay black in all orientations

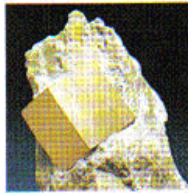
Such minerals are **isotropic**



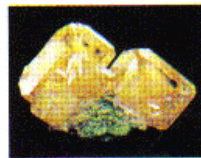
# Types of crystals

Table 4-1 Crystal Systems

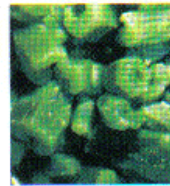
## Examples



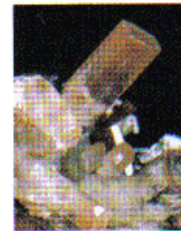
Pyrite



Wulfenite



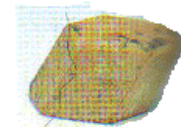
Pyromorphite



Topaz

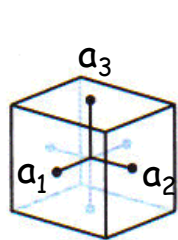


Gypsum

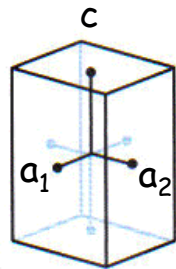


Feldspar

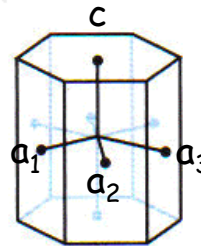
## Systems



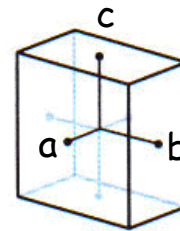
Cubic



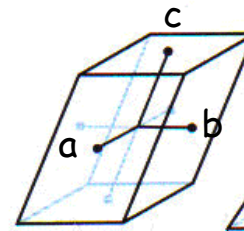
Tetragonal



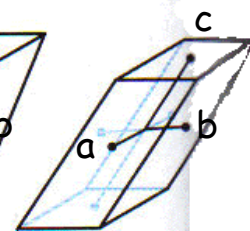
Hexagonal



Orthorhombic



Monoclinic



Triclinic

uniaxial  
Isotropic

biaxial

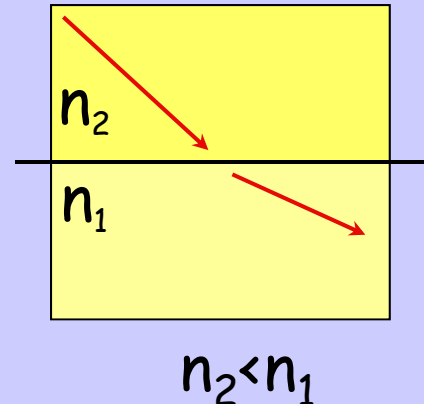
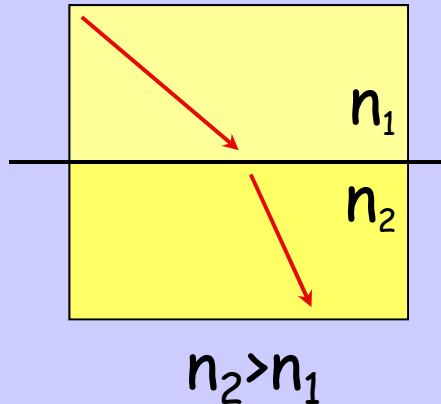
Anisotropic

triaxial

# Mineral properties: Index of refraction (R.I. or n)

$$n = \frac{\text{velocity in air}}{\text{velocity in mineral}}$$

Light is **refracted** when it passes from one substance to another; refraction is accompanied by a change in **velocity**

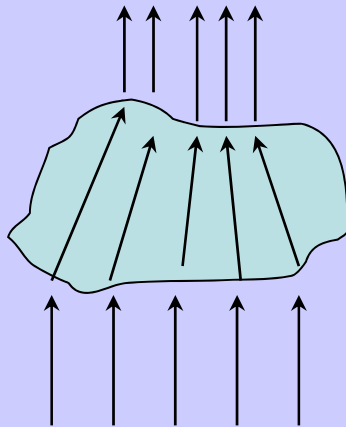


- n is a function of **crystallographic orientation** in anisotropic minerals
  - ⇒ **isotropic** minerals: characterized by **one** RI
  - ⇒ **uniaxial** minerals: characterized by **two** RI
  - ⇒ **biaxial** minerals: characterized by **three** RI
- n gives rise to 2 easily measured parameters: **relief** & **birefringence**

## What causes relief?

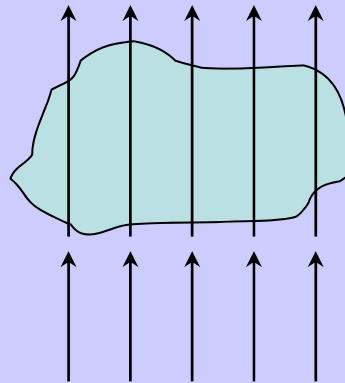
Difference in speed of light ( $n$ ) in different materials causes refraction of light rays, which can lead to focusing or defocusing of grain edges relative to their surroundings

Hi relief (+)



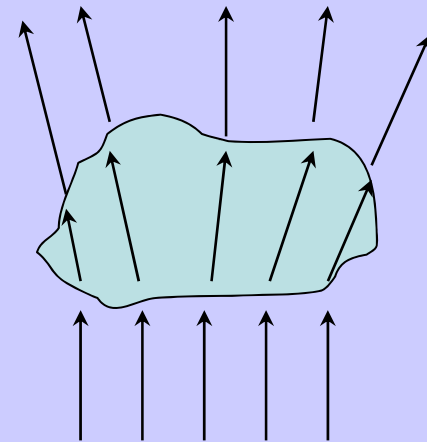
$$n_{xtl} > n_{epoxy}$$

Lo relief (+)



$$n_{xtl} = n_{epoxy}$$

Hi relief (-)



$$n_{xtl} < n_{epoxy}$$



# Mineral properties: **relief**

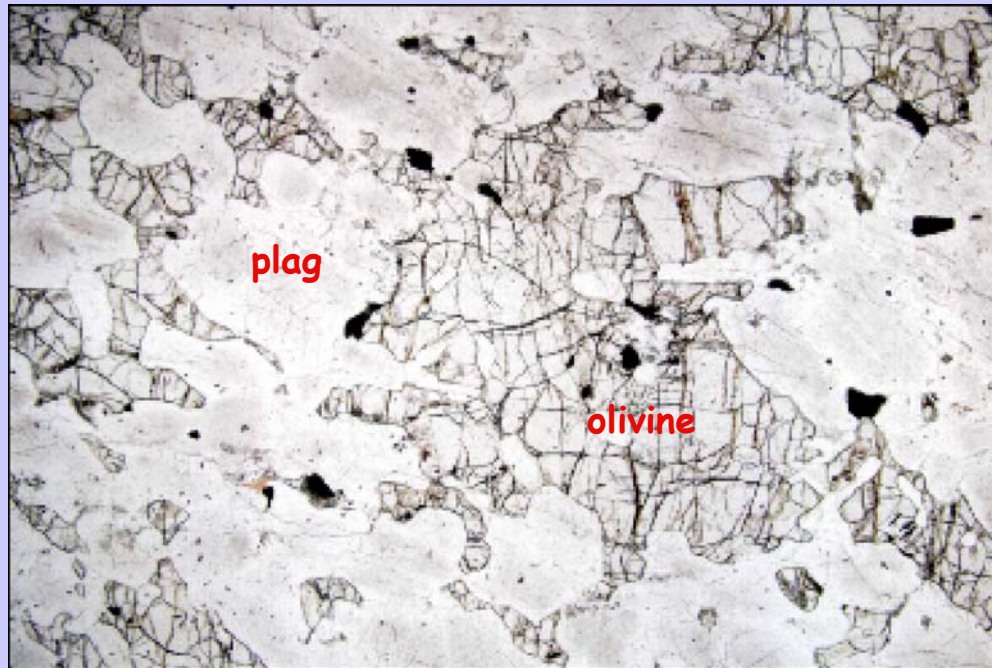
- Relief is a measure of the **relative difference in  $n$**  between a mineral grain and its surroundings
- Relief is determined visually, in **PPL**
- Relief is used to **estimate  $n$**



garnet:	$n = 1.72-1.89$
quartz:	$n = 1.54-1.55$
epoxy:	$n = 1.54$

# Mineral properties: **relief**

- Relief is a measure of the **relative difference in  $n$**  between a mineral grain and its surroundings
- Relief is determined visually, in **PPL**
- Relief is used to **estimate  $n$**

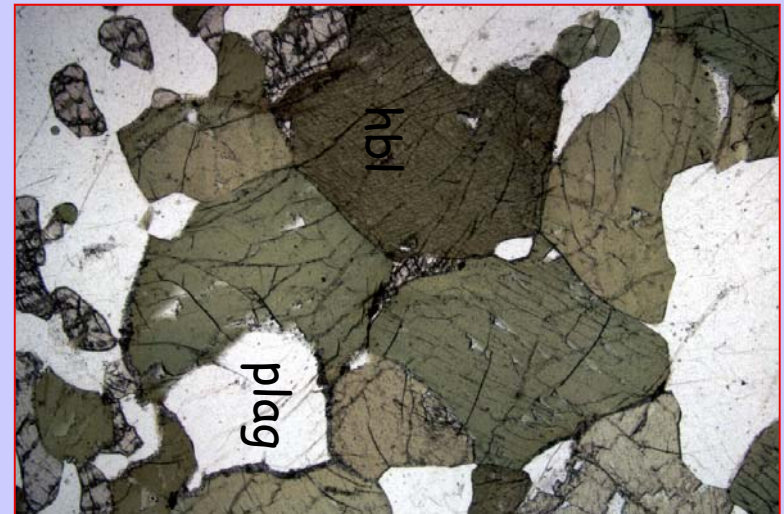
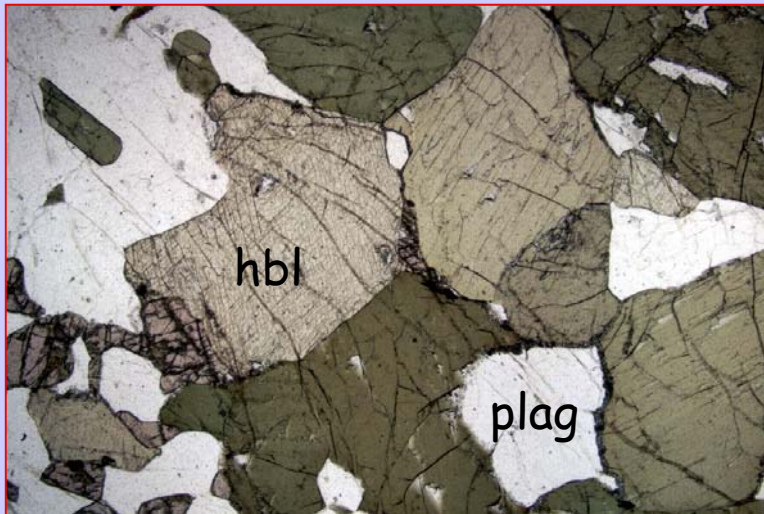


- Olivine has high relief
- Plag has low relief

olivine:	$n=1.64-1.88$
plag:	$n=1.53-1.57$
epoxy:	$n=1.54$

# Color and pleochroism

- **Color** is observed **only in PPL**
- Not an inherent property - changes with light type/intensity
- Results from selective absorption of certain  $\lambda$  of light
- **Pleochroism** results when different  $\lambda$  are absorbed differently by different crystallographic directions - rotate stage to observe

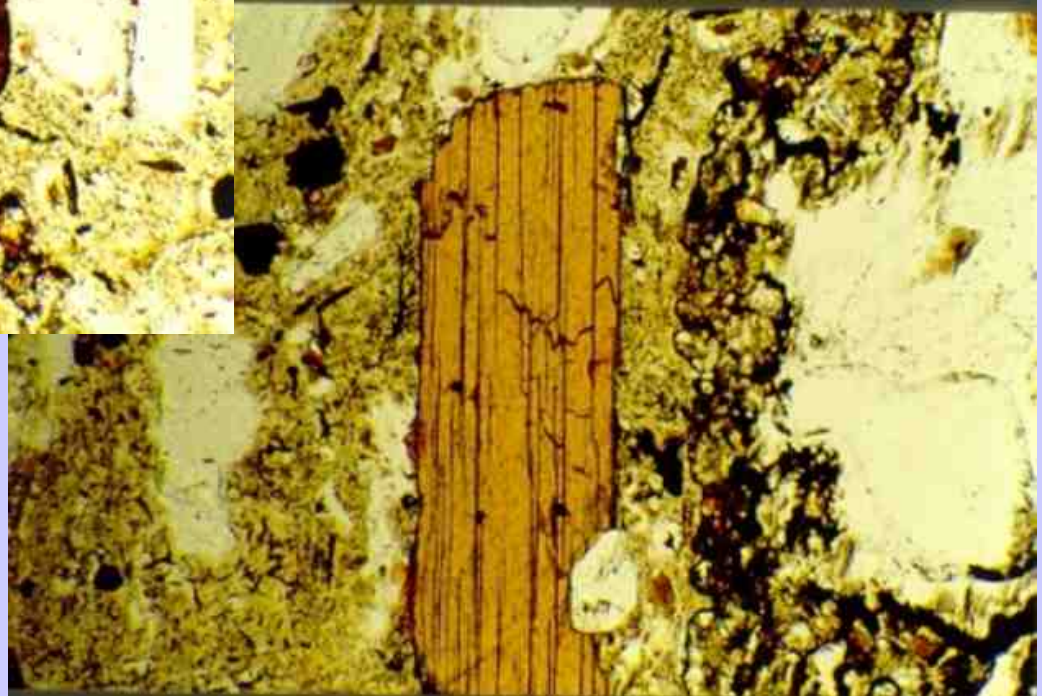
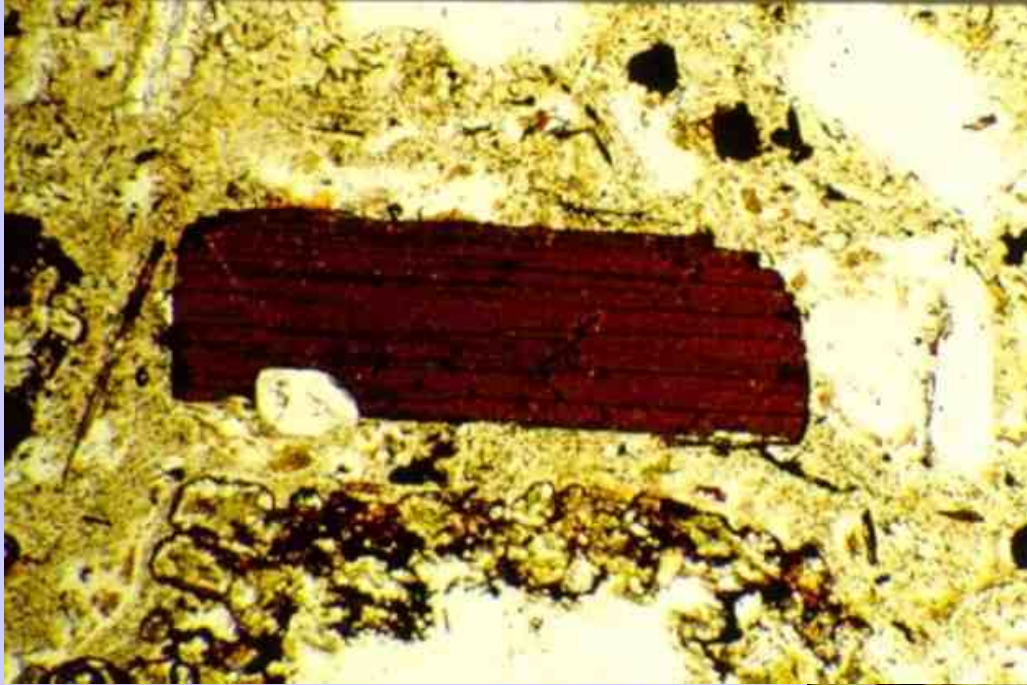


- Plagioclase is colorless
- Hornblende is pleochroic in olive greens

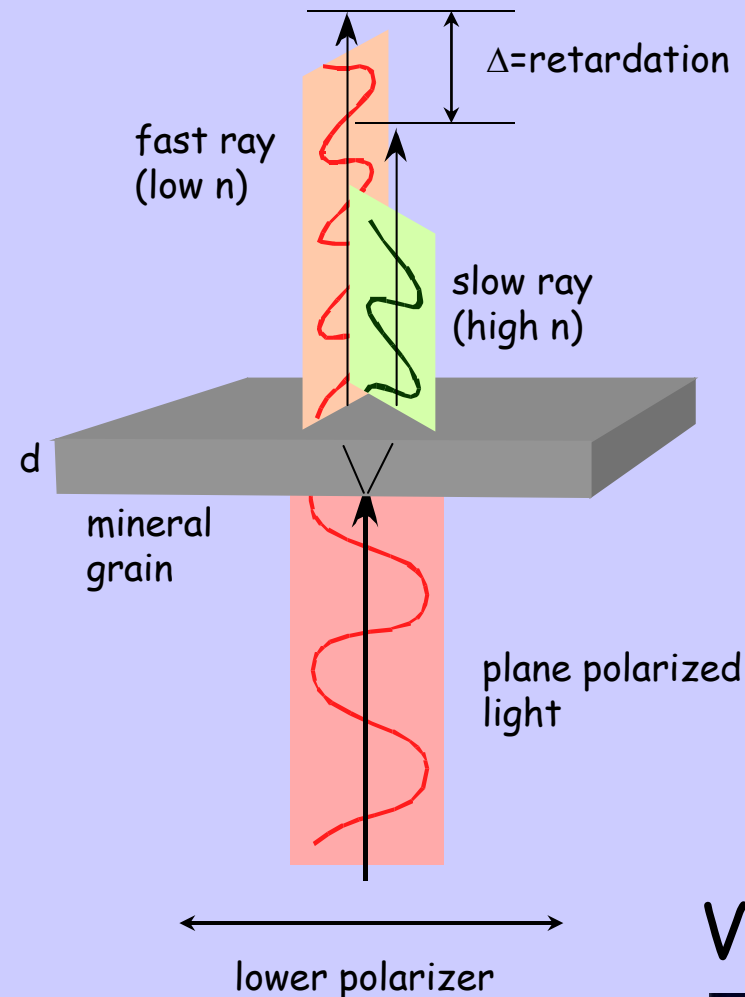


# Color and pleochroism

## Biotite



# Birefringence (possessing more than one index of refraction)/interference colors



Observation:

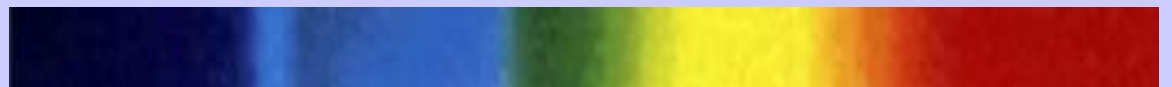
**frequency** of light remains unchanged during splitting, regardless of material

$$F = V/\lambda$$

if light speed changes,  $\lambda$  must also change

$\lambda$  is related to **color**; if  $\lambda$  changes, color also changes

Violet (400 nm) → Red (700 nm)



# Anisotropic crystals

## Calcite experiment and double refraction

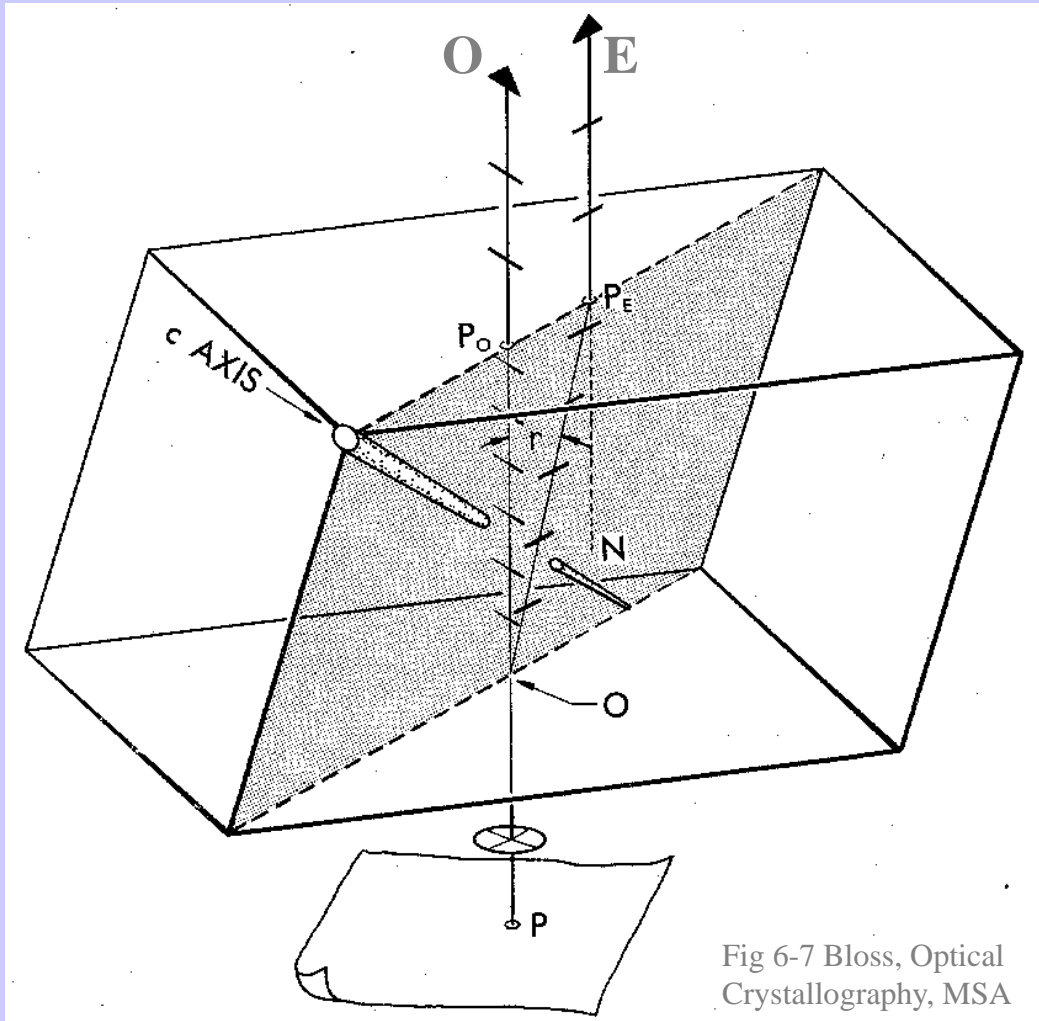


Fig 6-7 Bloss, Optical Crystallography, MSA

Double images

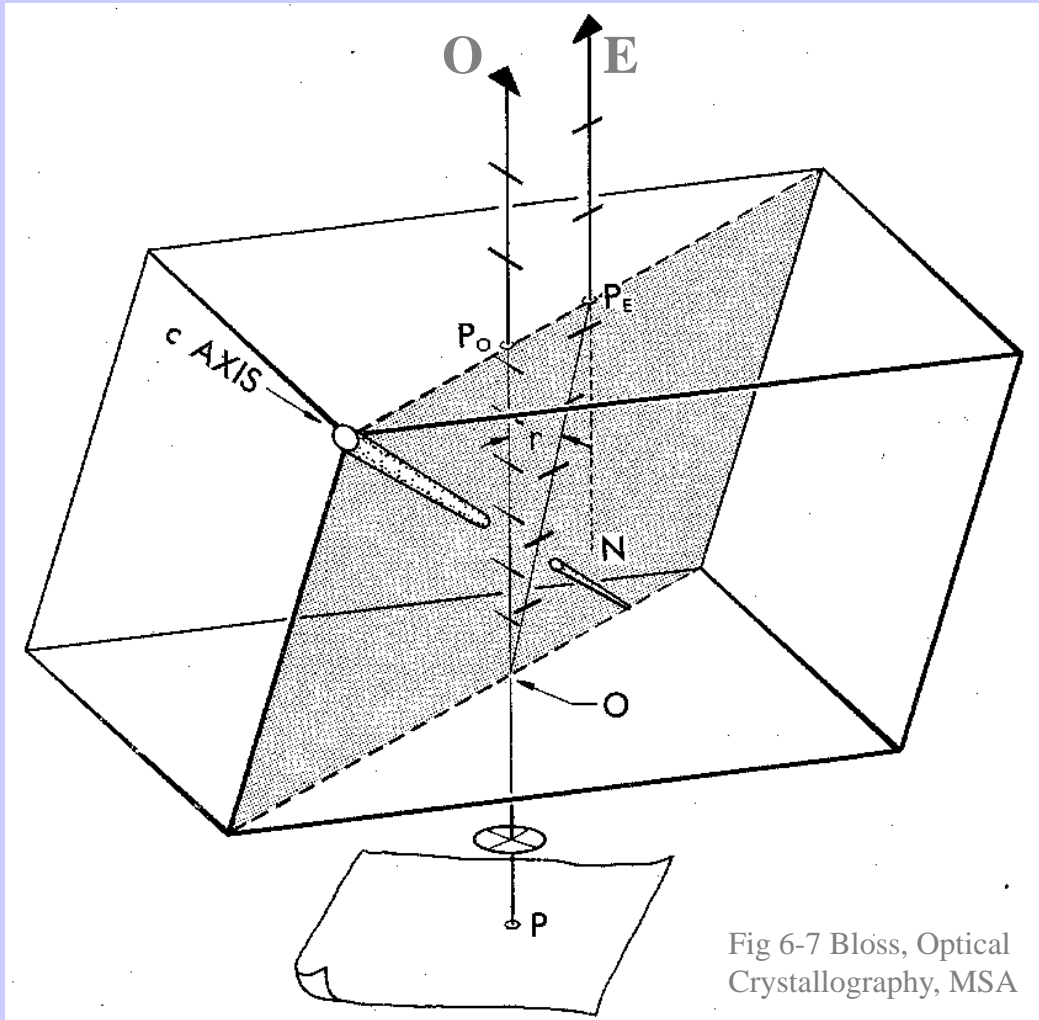
Ray → 2 rays with  
different propagation  
and vibration directions

Each is polarized ( $\perp$  each  
other)



# Anisotropic crystals

## Calcite experiment and double refraction



### O-ray (Ordinary)

Obeys Snell's Law and goes straight

Vibrates  $\perp$  plane containing ray and c-axis (“optic axis”)

### E-ray (Extraordinary)

Deflected

Vibrates in plane containing ray and c-axis

Fig 6-7 Bloss, Optical Crystallography, MSA

# Interference phenomena

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- Light waves may be **in phase** or **out of phase** when they exit xtl
- When out of phase, some component of light gets through upper polarizer and displays an **interference color**
- When one of the vibration directions is parallel to the lower polarizer, no light gets through the upper polarizer and the grain is "**at extinction**" (=black)  
For uniaxial crystals extinction is "parallel" to the polarizers, for biaxial crystals extinction is "inclined" to the polarizers.

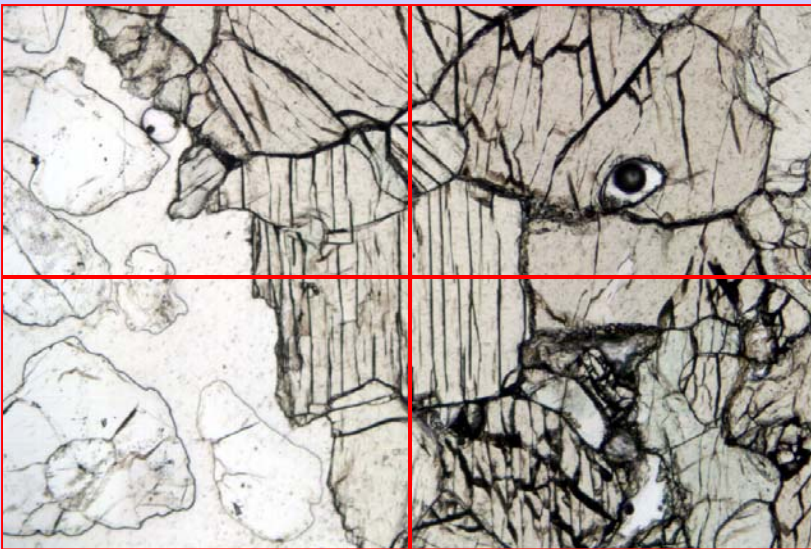
# Examples of extinction angle - parallel extinction

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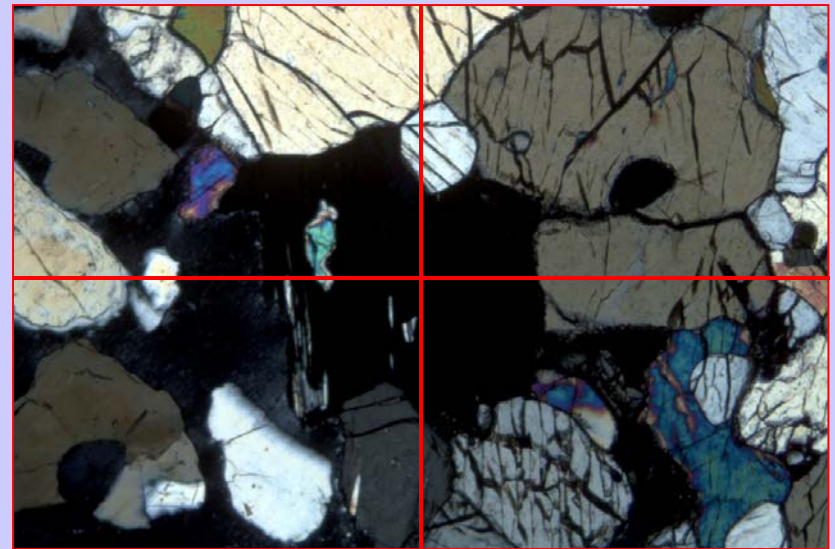
- All **uniaxial** minerals show **parallel extinction**
- **Orthorhombic** minerals show **parallel extinction**

*(this is because xtl axes and indicatrix axes coincide)*

orthopyroxene



PPL



XN



# Examples of extinction angle - inclined extinction

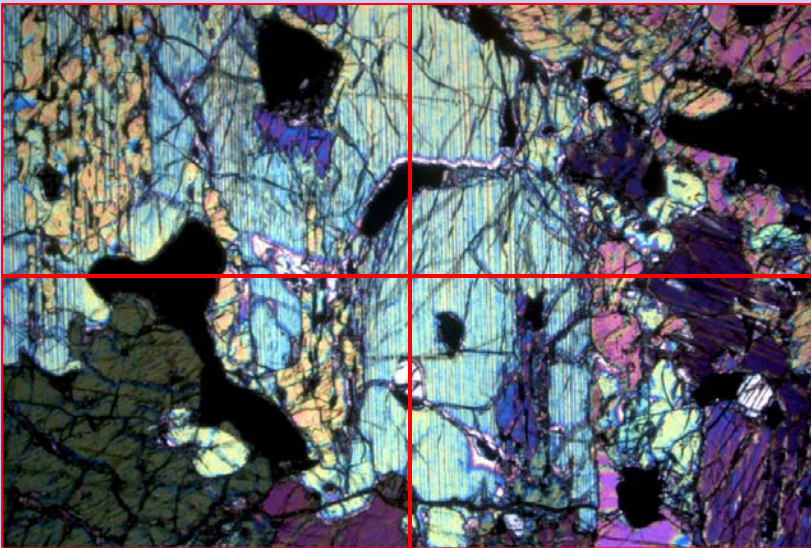
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**Monoclinic** and **triclinic** minerals:

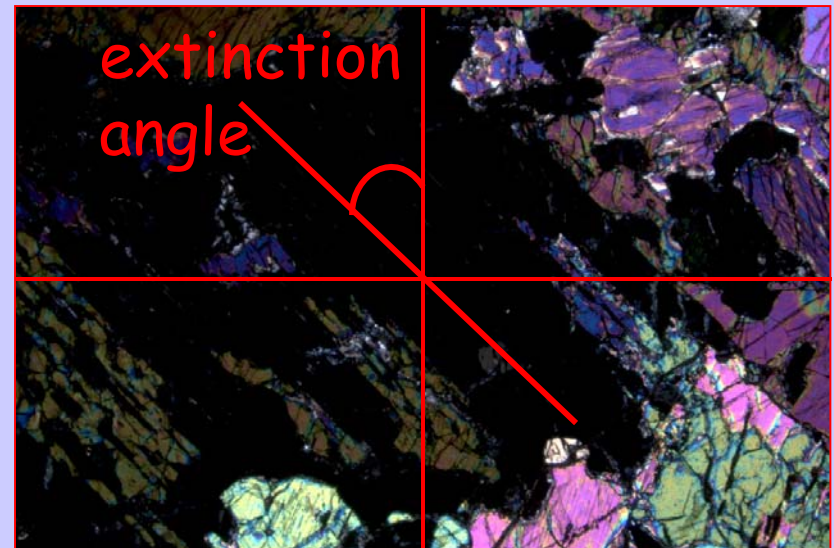
indicatrix axes do not coincide with crystallographic axes

*These minerals have **inclined extinction***

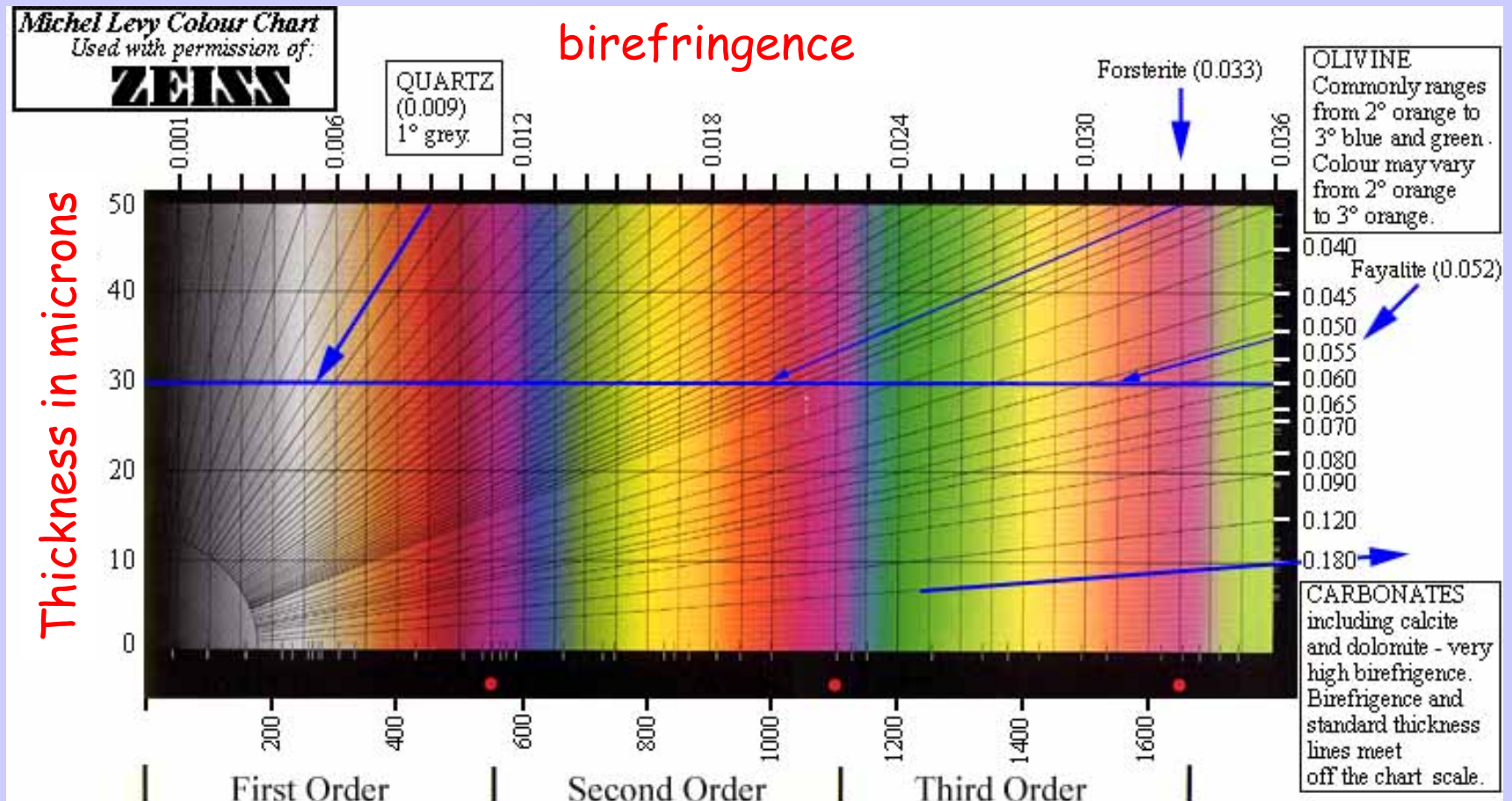
*(and extinction angle helps to identify them)*



clinopyroxene



# Birefringence/interference colors

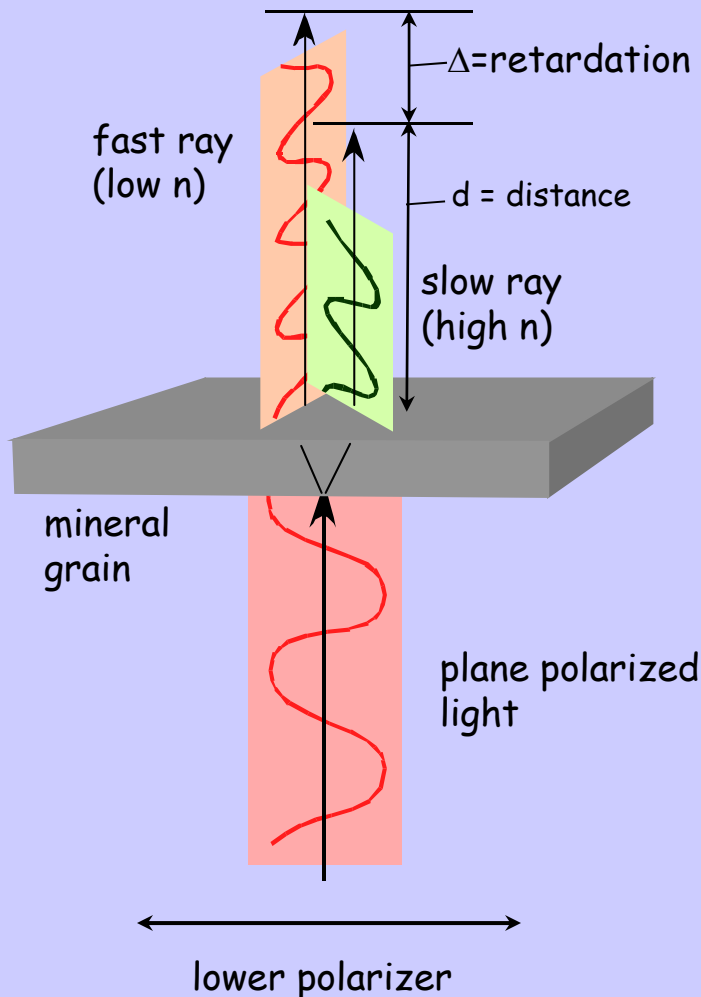


**Retardation in nanometers**

At time  $t$ , when slow ray 1<sup>st</sup> exits xtl:

Slow ray has traveled distance  $d$

Fast ray has traveled distance  $d + \Delta$



$$\text{time} = \text{distance}/\text{rate}$$

Slow ray:

$$t = d/V_{\text{slow}}$$

Fast ray:

$$t = d/V_{\text{fast}} + \Delta/V_{\text{air}}$$

Therefore:

$$d/V_{\text{slow}} = d/V_{\text{fast}} + \Delta/V_{\text{air}}$$

$$\Delta = d(V_{\text{air}}/V_{\text{slow}} - V_{\text{air}}/V_{\text{fast}})$$

$$\Delta = d(n_{\text{slow}} - n_{\text{fast}})$$

$$\Delta = d \delta$$

$\Delta$  = thickness of t.s. x birefringence



# Color chart

Shows the relationship between retardation, crystal thickness, and interference color

550  $\mu\text{m}$   $\rightarrow$  red violet

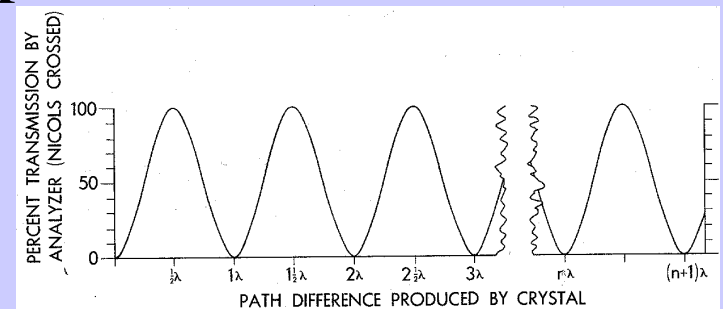
800  $\mu\text{m}$   $\rightarrow$  green

1100  $\mu\text{m}$   $\rightarrow$  red-violet again (note repeat  $\uparrow$ )

0-550  $\mu\text{m}$  = “1<sup>st</sup> order”      550-1100  $\mu\text{m}$  = 2<sup>nd</sup> order

1100-1650  $\mu\text{m}$  = 3<sup>rd</sup> order...

Higher orders are more pastel



# FYI: Miller Indices (definition)

the rules for Miller Indices are:

Determine the intercepts of the face along the crystallographic axes, *in terms of unit cell dimensions*.

Take the reciprocals

Clear fractions

Reduce to lowest terms

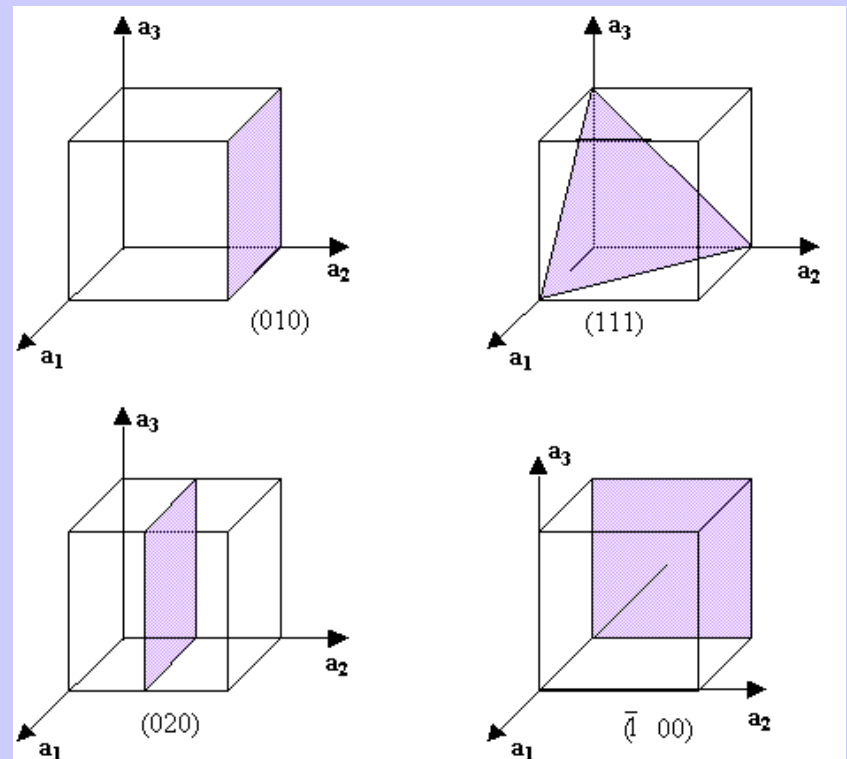


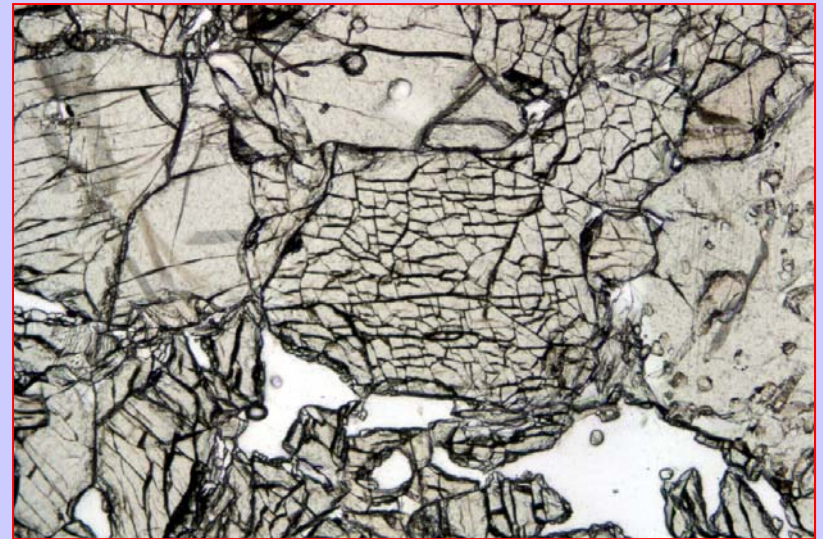
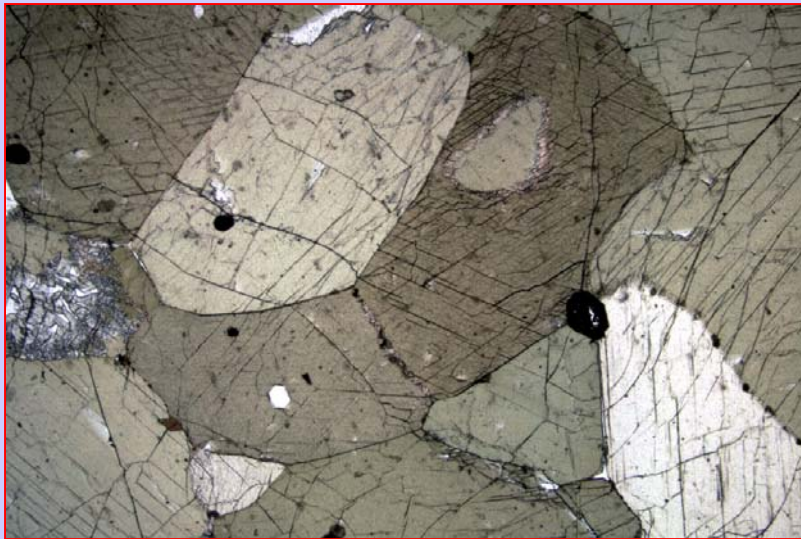
Figure 20: Miller indices of some planes in cubic crystal.

# FYI: Cleavage = the way a mineral preferentially breaks (reflects internal structure)

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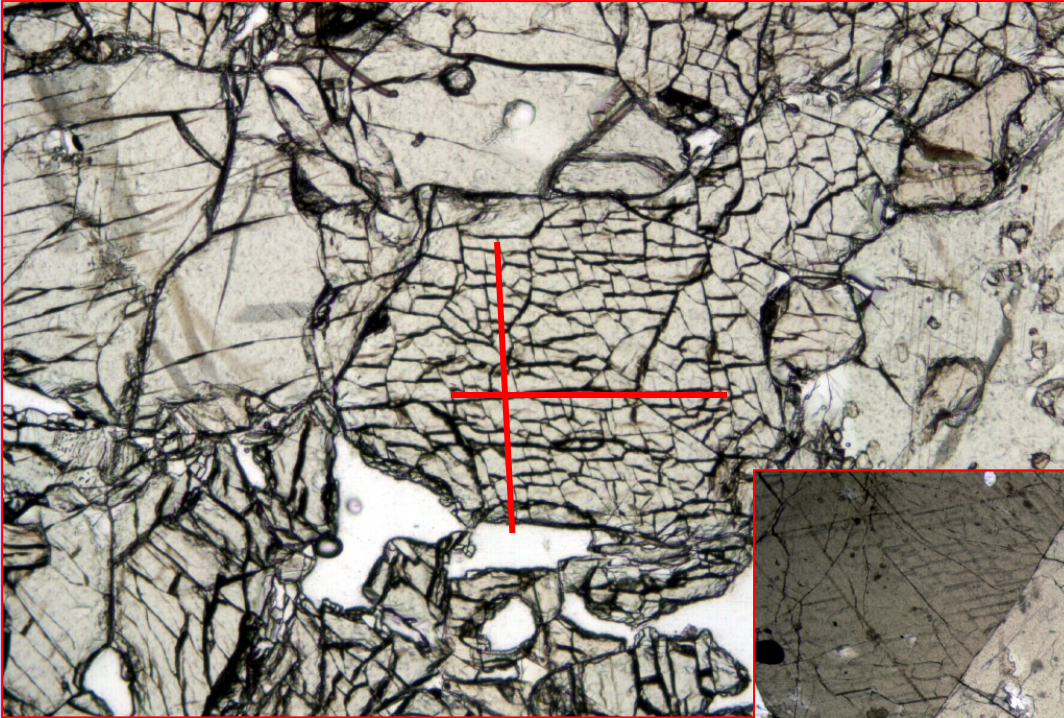
Most easily observed in PPL (upper polarizer out),  
but visible in XN as well

- No cleavages: quartz, olivine
- 1 good cleavage: micas
- 2 good cleavages: pyroxenes, amphiboles



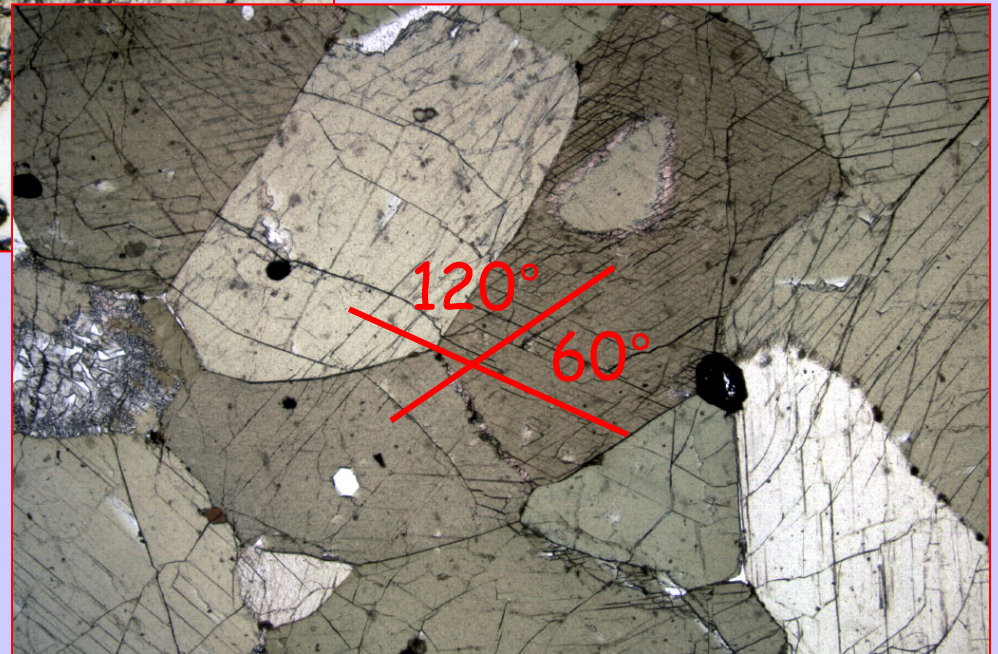


## FYI: Cleavage examples



2 cleavages  
intersecting  
at  $\sim 90^\circ$   
pyroxene

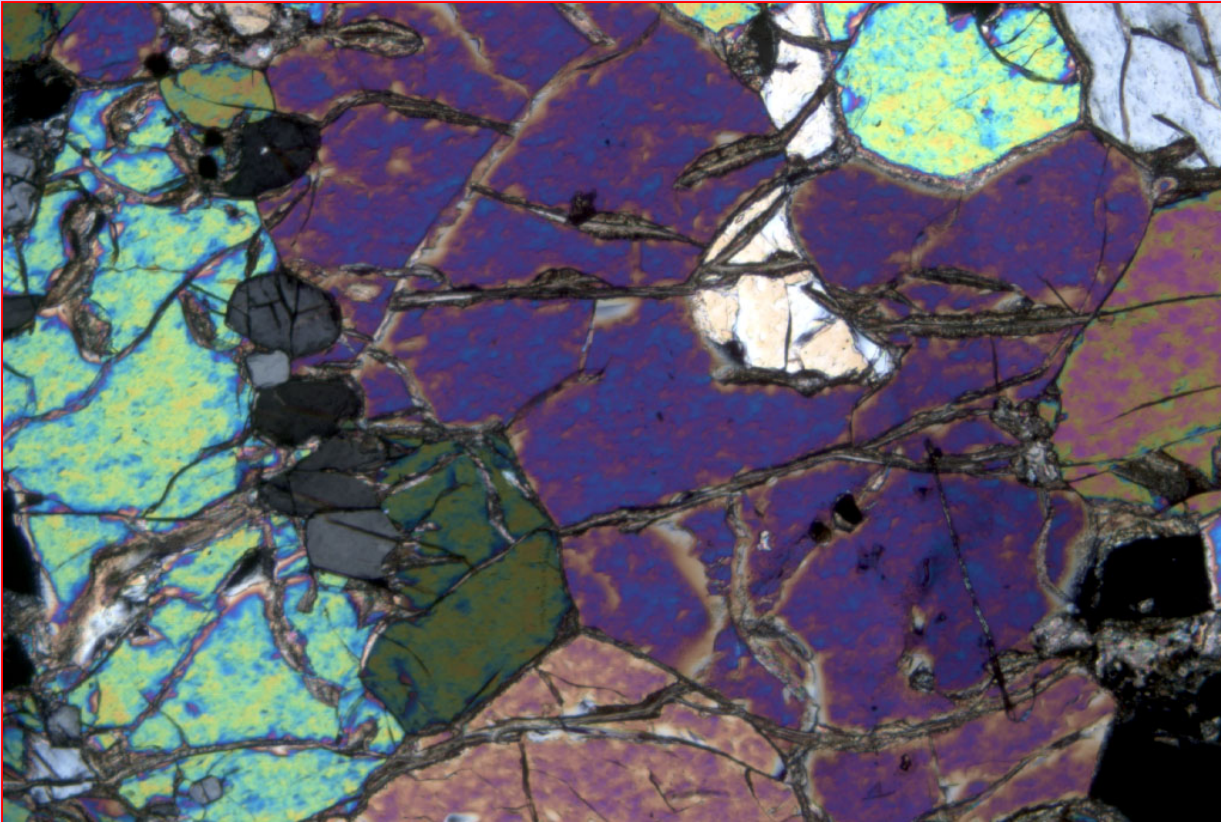
2 cleavages  
intersecting  
at  $60^\circ/120^\circ$ :  
amphibole





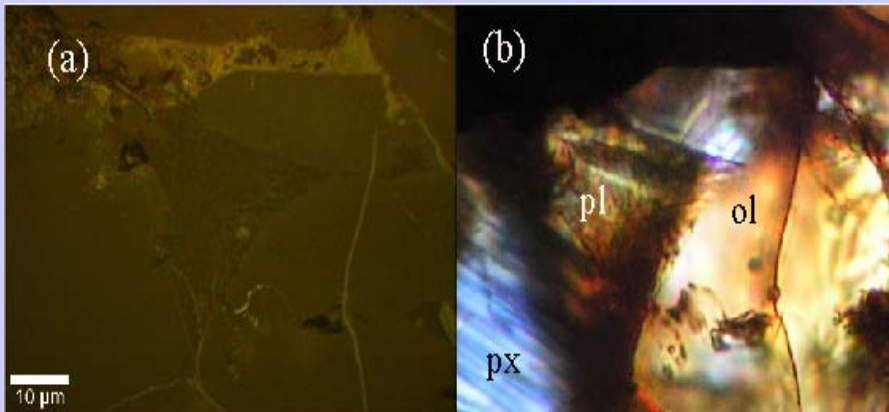
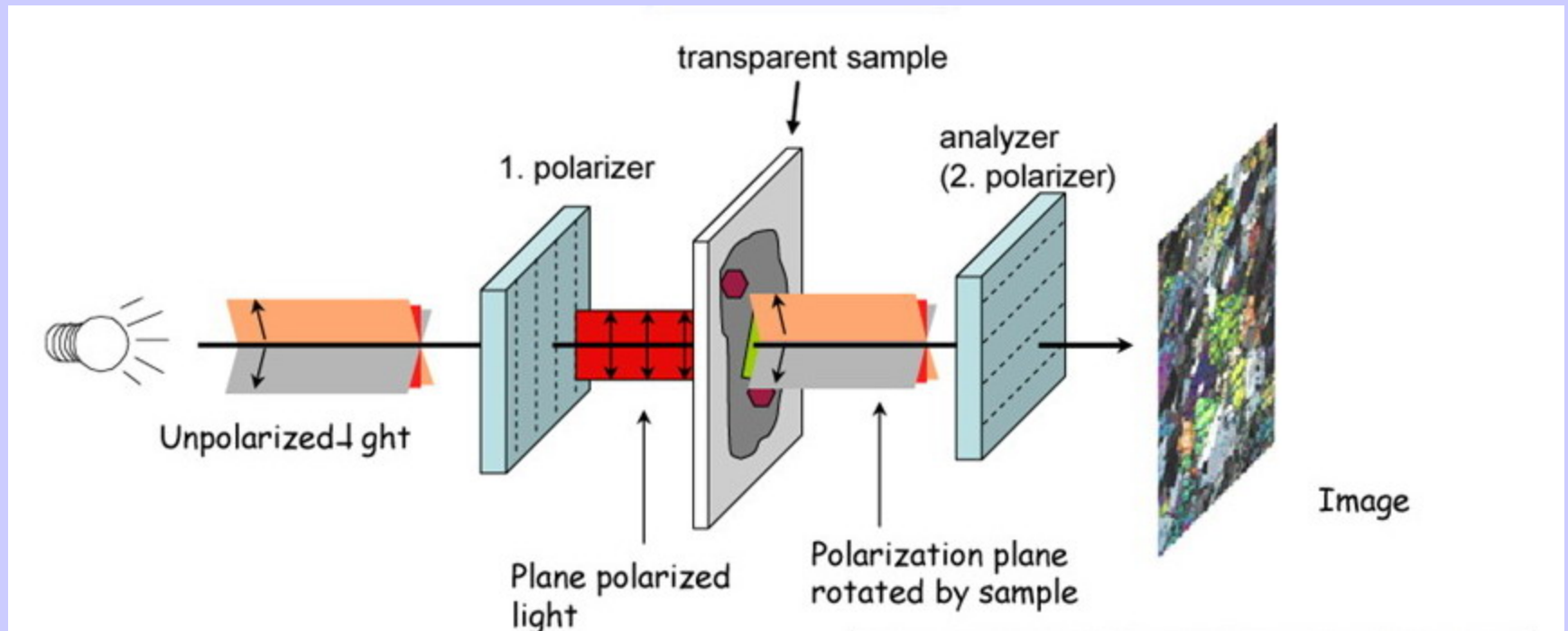
## FYI: Cleavage examples

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random fractures,  
no cleavage:  
olivine

# Polarized Light Microscopy



Reflected (a) and cross polarized transmitted (b) light images of RC 05: *ol* = olivine; *pl* = plagioclase; *px* = pyroxene.